

# Hurricanes and typhoons in the global climate system

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all PRIMAVERA partners (models and analysis)

*With many thanks to Suzana Camargo, Tom Knutson and Jim Kosssin*

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<sup>7</sup>ECMF

<sup>8</sup>ISAC-CNR

<sup>9</sup>Oxford University

<sup>10</sup>NCAS-CMS

# Motivation: TCs as rare, albeit significant contributors to climate

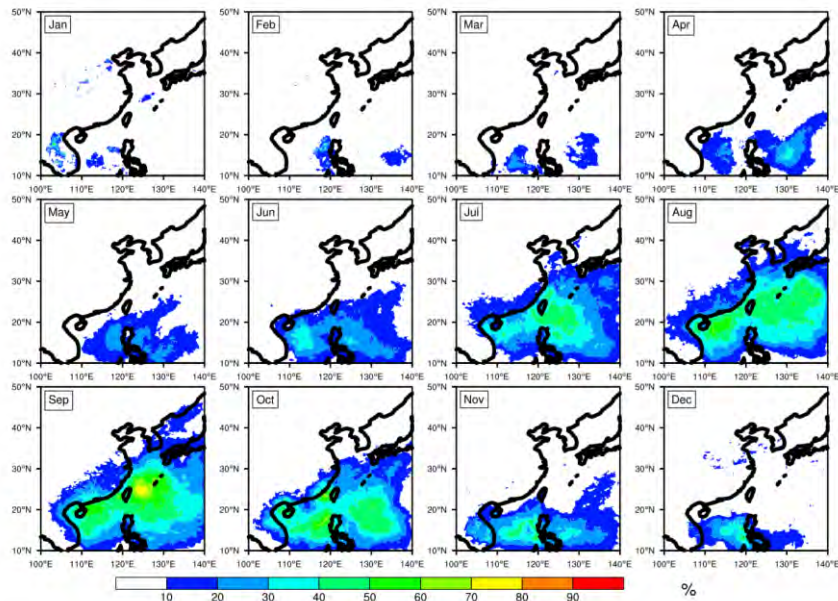
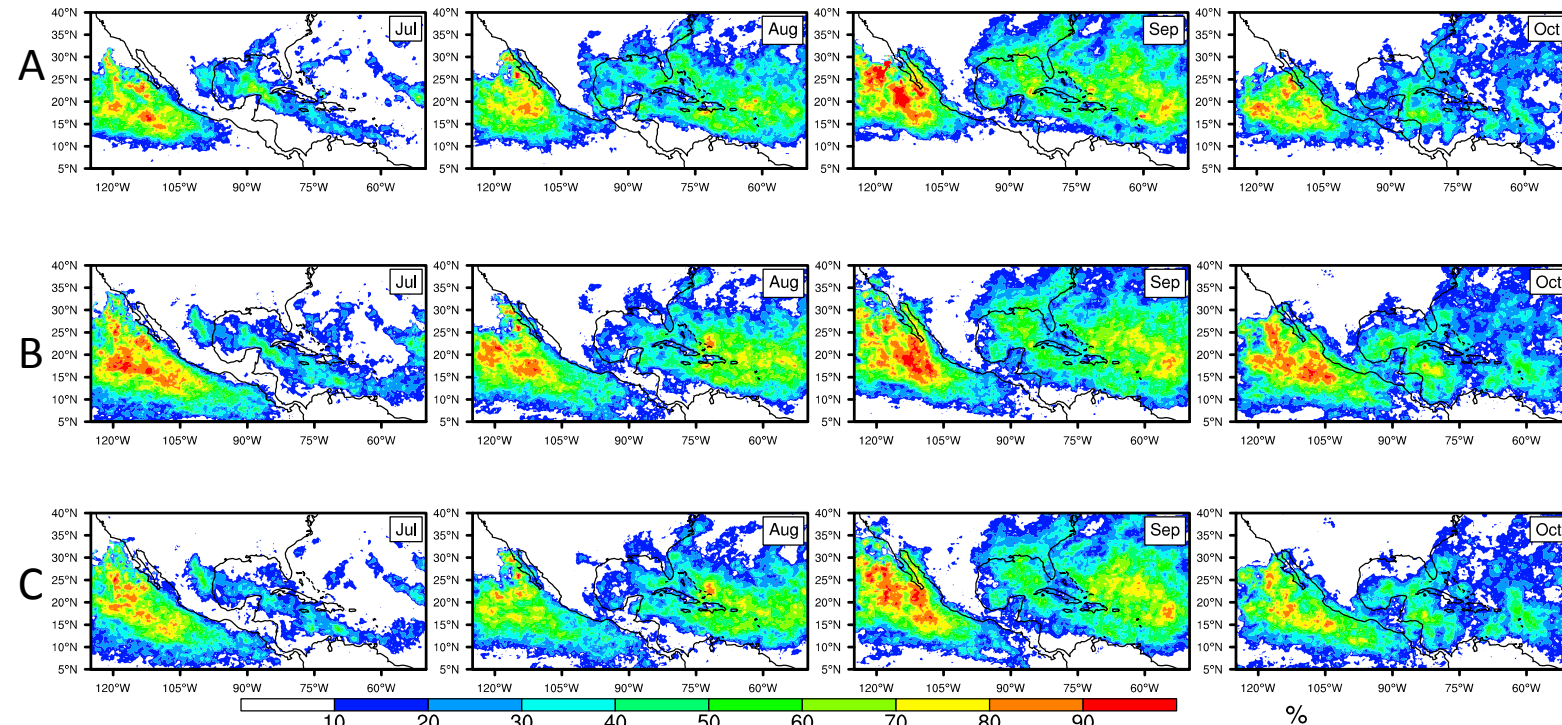


FIG. 2. Monthly mean fractional contribution of TC rainfall amount to the total rainfall calculated using TRMM 3B42 rainfall data. Units: %.



Contribution of TCs to the extreme rainfall (amount fraction) (%) from July to October, employing TC tracks from (a) IBTrACS, (b) JRA-55 and (c) ERA-Interim. Climatology for 1998-2015

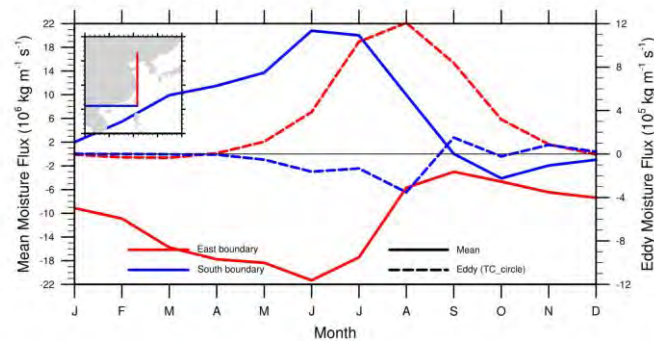


FIG. 5. Seasonal cycle of monthly mean vertically integrated moisture flux passing through the southern (blue) and eastern (red) boundaries. Mean flow moisture fluxes are shown as solid lines and TC eddy moisture fluxes as dash lines. The inner panel shows the definition of the southern and eastern boundaries. Units: kg/m/s.

Guo et al. 2017

**Re-analyses very likely under-estimating the role of TCs in producing precipitation and moisture transports.**

What is the role of model resolution, model physics, initialisation (Data Assimilation)?

Franco-Diaz et al.  
2019, submitted to  
Clim Dyn

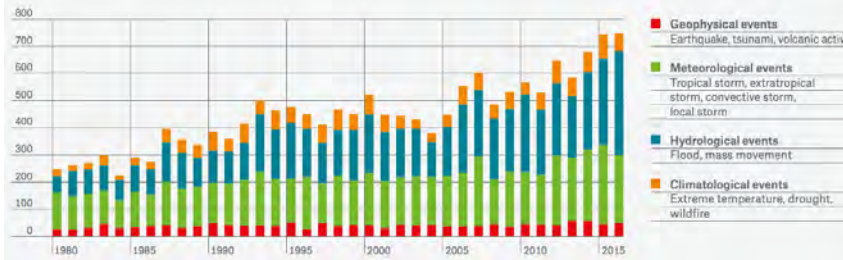
# Recent natural catastrophes: comparing 2011 with other years

NatCatSERVICE

Loss events worldwide 2017 ~300 US\$ billion  
Geographical overview

Munich RE

Number of loss events 1980-2016

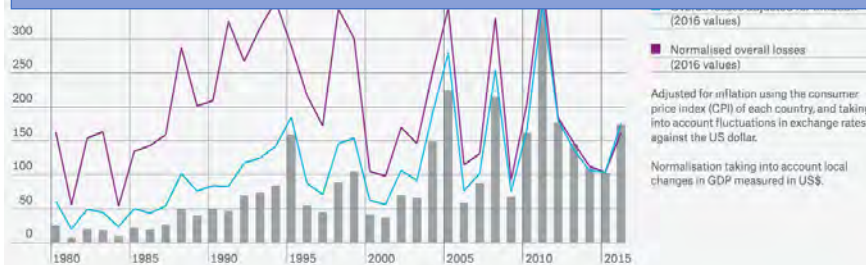


Overall losses and insured losses 1980-2016 (in US\$ bn)

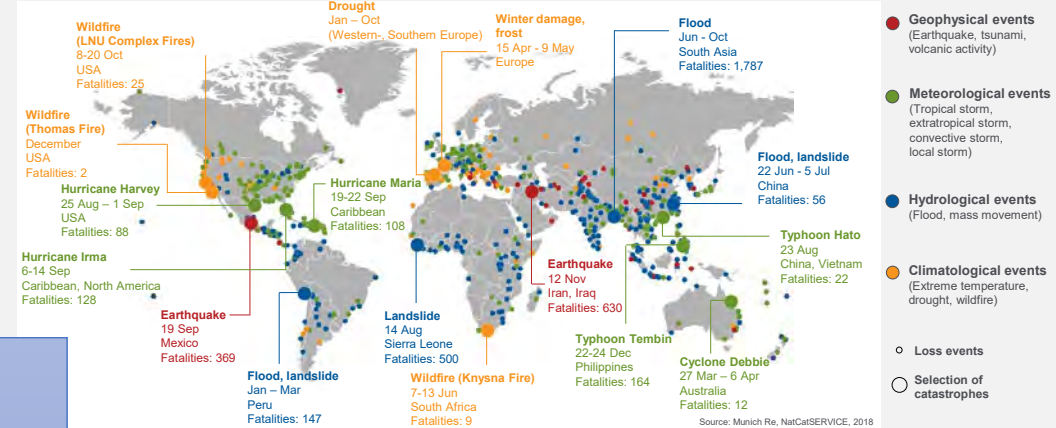
## Overseas Development work

Global annual loss is:

- Order of 100 billion US\$ each year
- Mostly HydroMet
- Uninsured (2/3)**
- Often governed by non-local processes
- Located in developing countries, where insurance cannot and will not operate, because there is no suitable evidence base.



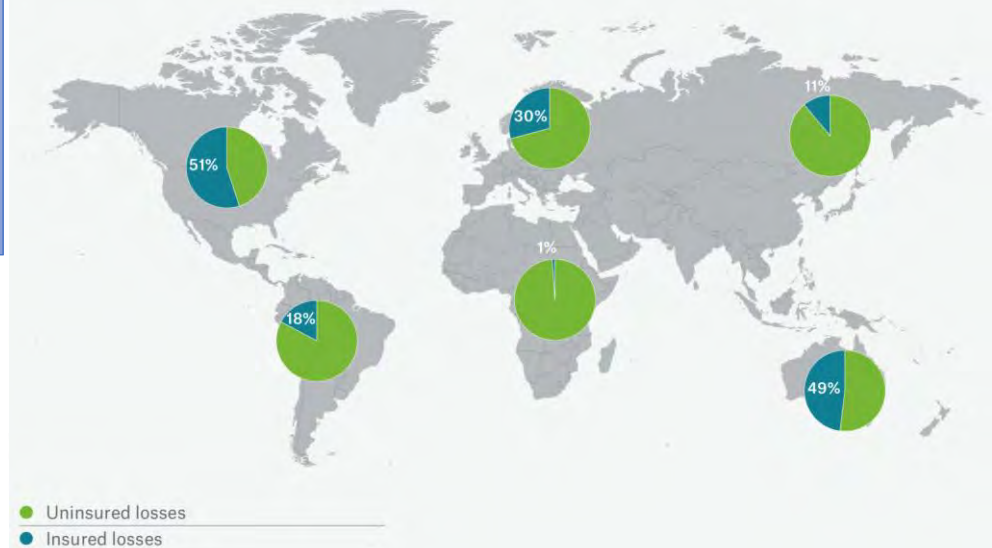
Source: Munich Re NatCatSERVICE



Ünchener Rückversicherungs-Gesellschaft, NatCatSERVICE - As at January 2018

Loss events 2016

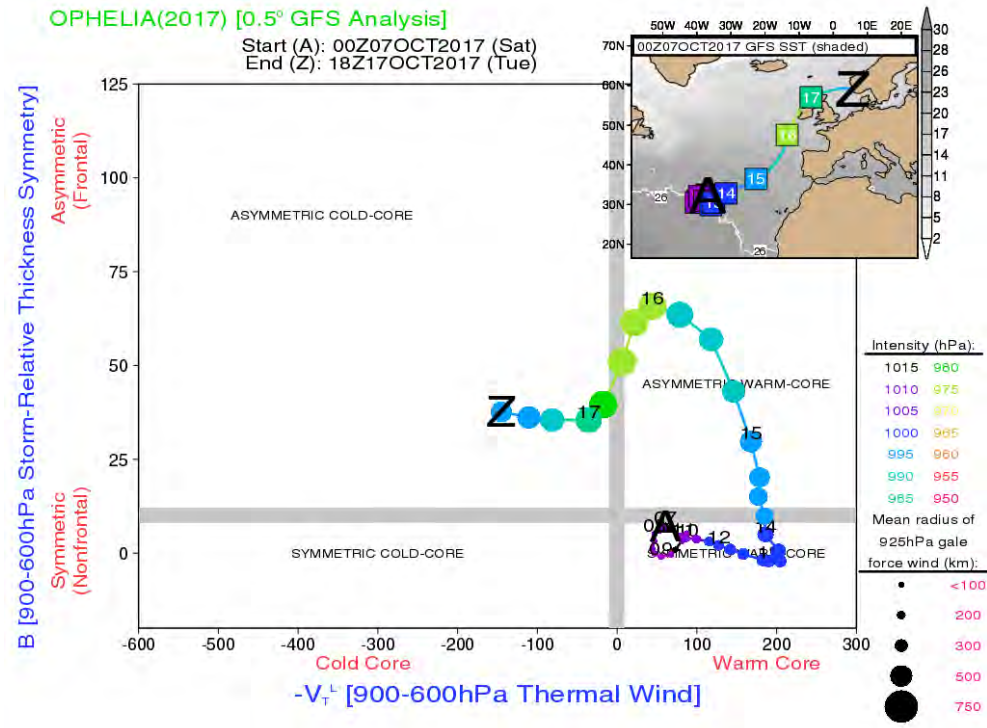
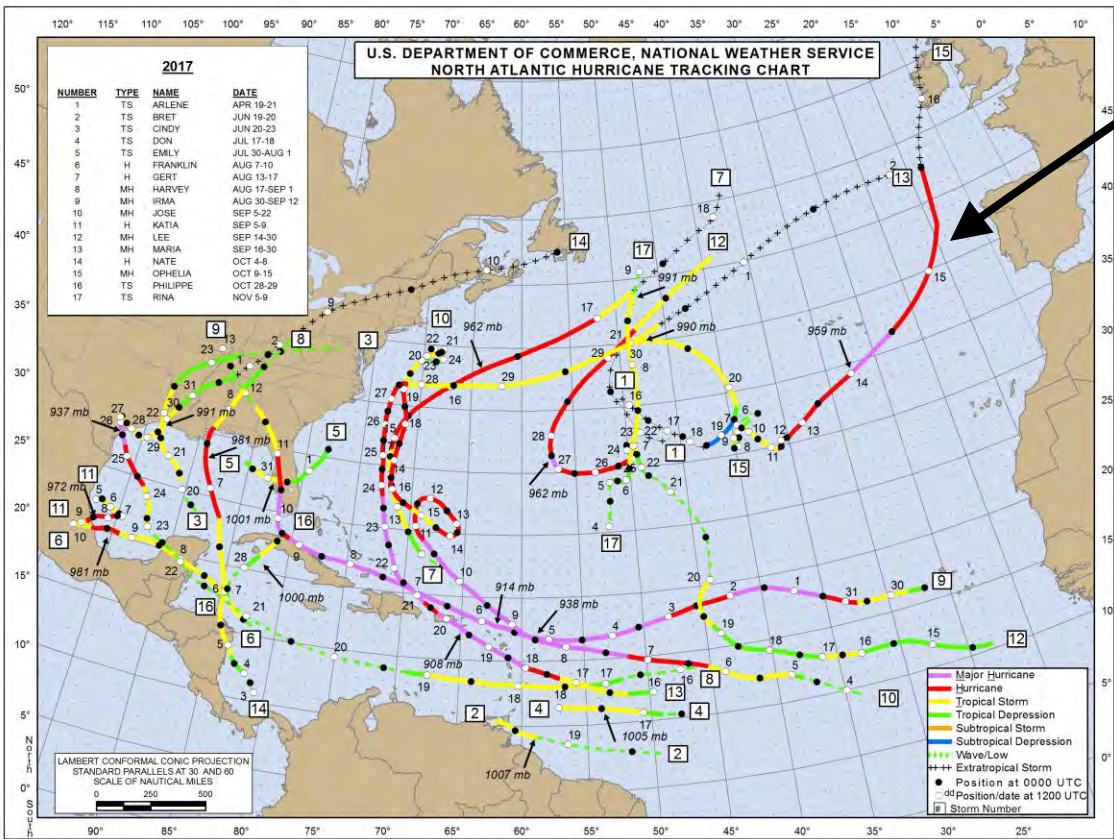
Insured losses as a percentage of overall losses for each continent



Munich Re 2011-2017

# Europe example: ex-hurricane Ophelia (2017)

The easternmost Atlantic hurricane on record caused three fatalities and \$65 million insured damages.



Cornwall, south-western UK (BBC)



County Kerry, Ireland (Irish Examiner)

Tropical-to-extratropical transition of cyclones exposes mid-latitude regions to hurricane-type hazards.

# Early GCM studies: 70s-80s

## *“Hurricane-type vortices”*

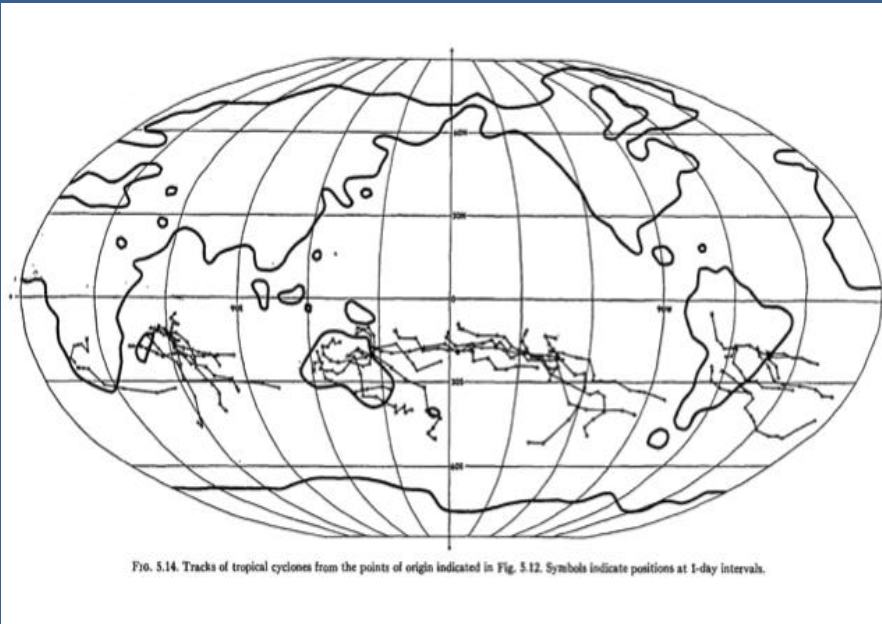
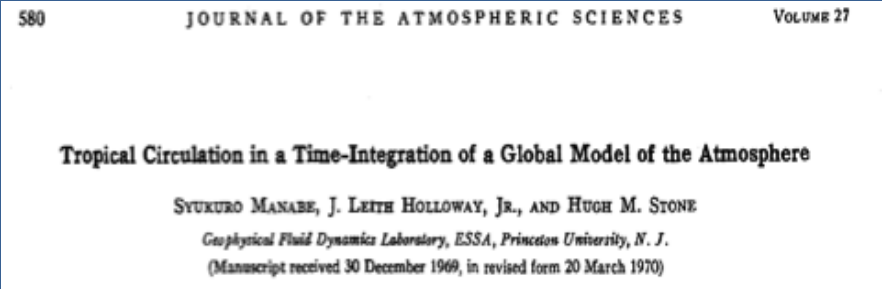
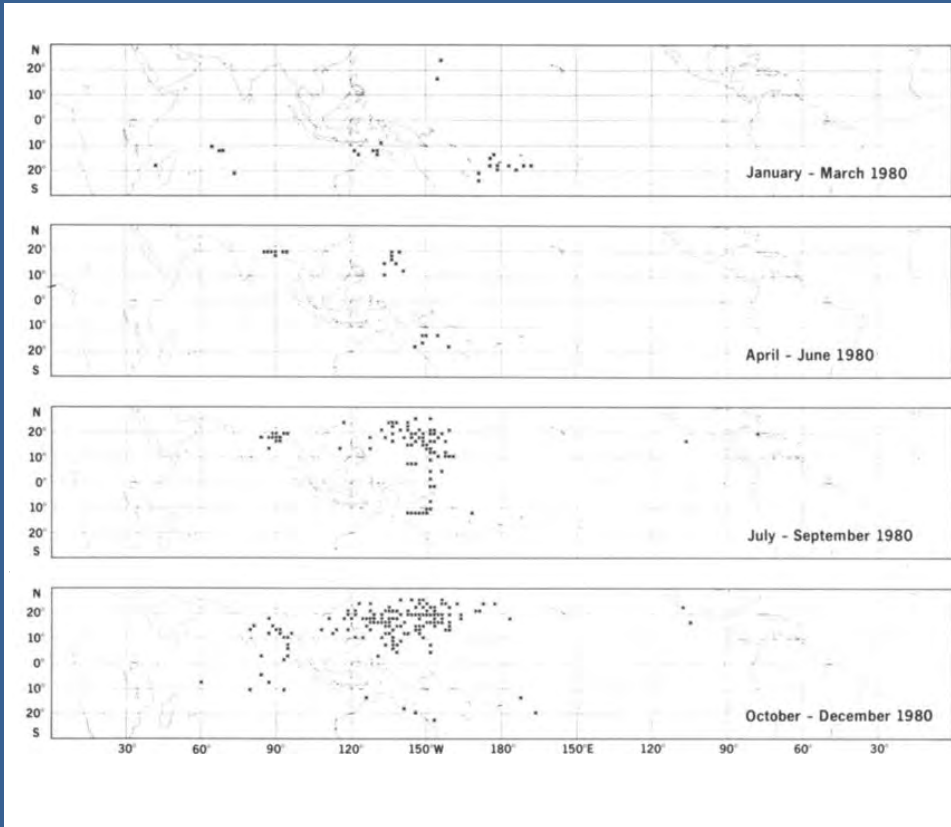


FIG. 5.14. Tracks of tropical cyclones from the points of origin indicated in Fig. 5.12. Symbols indicate positions at 1-day intervals.

This is not in anyway a surprise, since the horizontal and vertical resolutions which are used in large-scale numerical models cannot satisfactorily describe the small-scale features of these phenomena. However, tropical cyclones have occasionally been observed to have dimensions over 2000 km, in particular in the western North Pacific. See for example the study of the super typhoon Tip (Dunnavan and Diercks, 1980). A numerical model with a resolution of  $3^\circ$  or less would therefore have the potential possibility of reproducing a vortex of this dimension, although it

basis carried out 10-day global forecasts 5 times a week from August 1, 1979 and daily from August 1, 1980. The following study is concerned with the period January 1, 1980–December 31, 1980. The numerical model has at times generated vortices with a structure and behaviour resembling intense tropical cyclones. The developments have mostly taken place around days 4 and 5 of the forecasts. Moreover, they have been found in areas and at times where tropical cyclones normally occur and they have shown a similar sensitivity to the sea-surface temperature as has been found in



# Climate change and TCs

changes may occur as a result of global warming. Models suggest that a doubling of atmospheric CO<sub>2</sub> tropical sea surface temperatures (SSTs) by as much as 2°C, so this speculation would appear to have some basis.

Manabe et al. (1987) attempted to examine this issue quantitatively by modeling the tropical cyclone as a Carnot heat engine in which heat is put into the atmosphere at the temperature of the sea surface. In an idealized model, he suggested that the warmer SSTs in a CO<sub>2</sub> world would increase the maximum sustainable intensity of tropical cyclones. He noted, however, that his analysis was based on a number of simplifying assumptions.

Manabe et al. (1987) employed terrain-following sigma coordinates. Solar radiation at the top of the atmosphere varies seasonally but not diurnally. Surface temperatures for land points are computed from a heat balance assuming no heat storage in the ground, and both snow cover and soil moisture are predicted. The moist convective adjustment scheme of Manabe et al. (1965) is used to parameterize convection.

Two versions of the model were used: a low resolution version with spherical harmonics truncated rhomboidally at wavenumber 15 (R15) and a high resolution version with spherical harmonics truncated rhomboidally at wavenumber 30 (R30). The latitude-longitude spacing of the control grids are 4.5° by 7.5° and 2.25° by 3.75°, respectively.

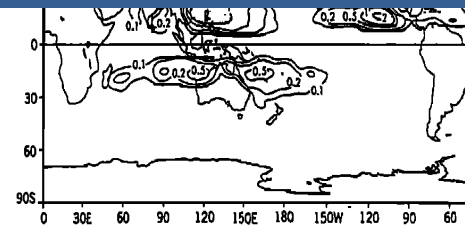
the atmosphere, where low-level vorticity, low-level shear, and other dynamical parameters have a large influence on the climatology (Gray, 1968, 1982). McBride et al. (1987) demonstrated that the antecedent conditions, and the structure of the developing disturbance in the model, were influenced by divergence rather than vorticity once again by a different physical mechanism to that produced in the atmosphere. This strong reliance on an idealized model may be a major limitation of the study.

vortices (and "storm days") than doubling of carbon dioxide. With this level of uncertainty in the control run, it is difficult to say that current climate GCMs are appropriate tools for exploring the relationship between greenhouse warming and tropical storm activity certainly seem to be overcautious.

In summary, there is no reason why GCMs should not be able to simulate the tropical atmosphere and, ideally, the tropical cyclone climatology. However, the current state of the art is such that the results are highly uncertain.

Manabe et al. (1987) would probably require substantial extrapolations beyond the intended scope of BM90.

Manabe et al. (1987) is critical of our use of a six-month season for examining tropical storm activity. Its use in BM90 resulted from our desire to develop an objective, automated technique for distinguishing tropical storms. As a first attempt, we assumed all storms identified within a specified spatial and temporal domain were tropical storms.

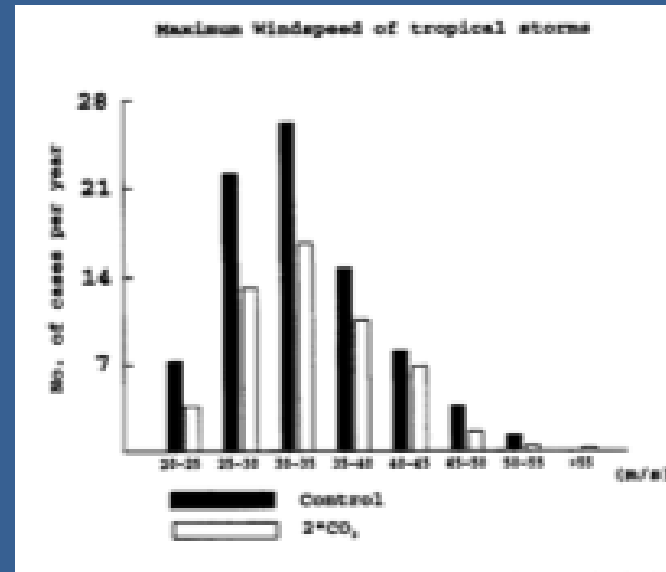


*Tellus* (1996), 48A, 57-73  
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TELLUS  
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## Will greenhouse gas-induced warming over the next 50 years lead to higher frequency and greater intensity of hurricanes?

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# TC seasonal variability

APRIL 1997

VITART ET AL.

745

## Simulation of Interannual Variability of Tropical Storm Frequency in an Ensemble of GCM Integrations

E. VITART

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J. L. ANDERSON AND W. F. STERN

*GFDL/NOAA, Princeton University, Princeton, New Jersey*

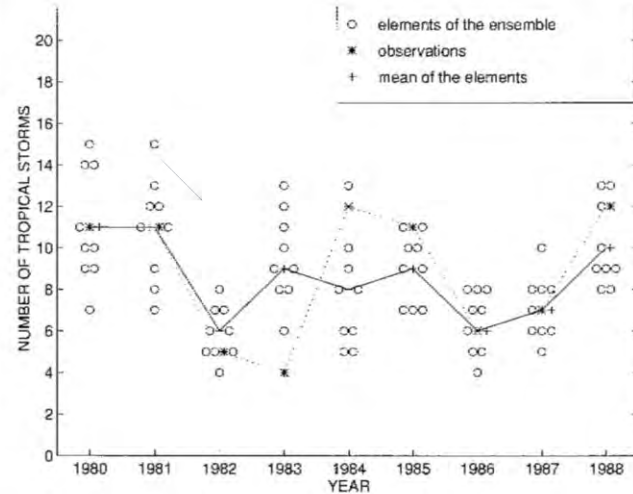


FIG. 7. Ensemble simulation of number of tropical storms and observed number of tropical storms over the western North Atlantic. The solid line represents observed tropical storm numbers. The dotted line represents the mean tropical storm numbers of the elements of the ensemble. Each circle represents the tropical storm number of one element of the ensemble.

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WEATHER AND FORECASTING

VOLUME 17

## Improving the Detection and Tracking of Tropical Cyclones in Atmospheric General Circulation Models

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*International Research Institute for Climate Prediction, Lamont-Doherty Earth Observatory of Columbia University, Palisades, New York*

(Manuscript received 25 February 2002, in final form 10 July 2002)

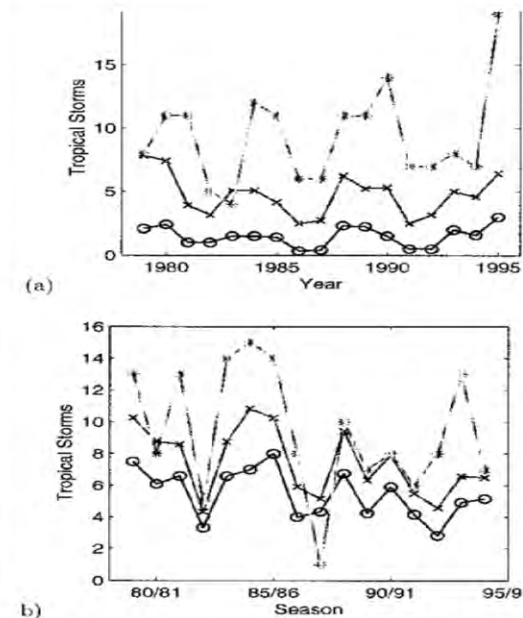
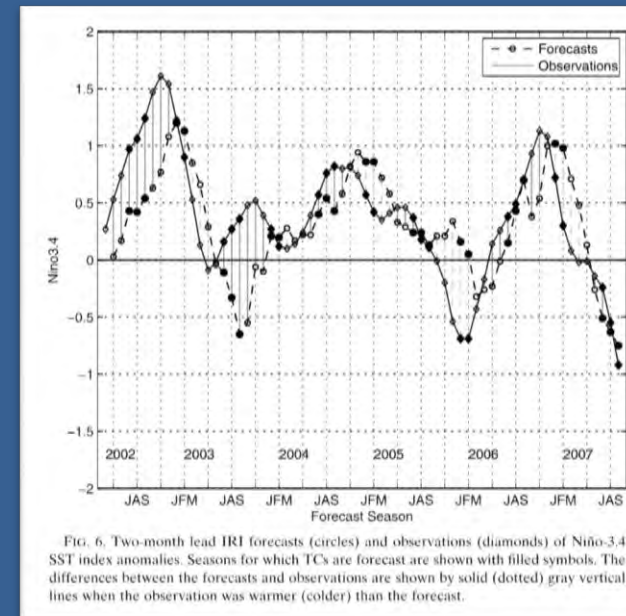
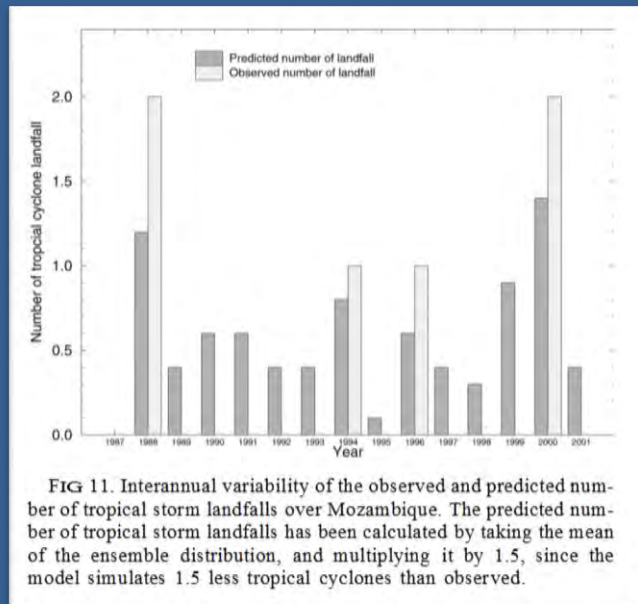
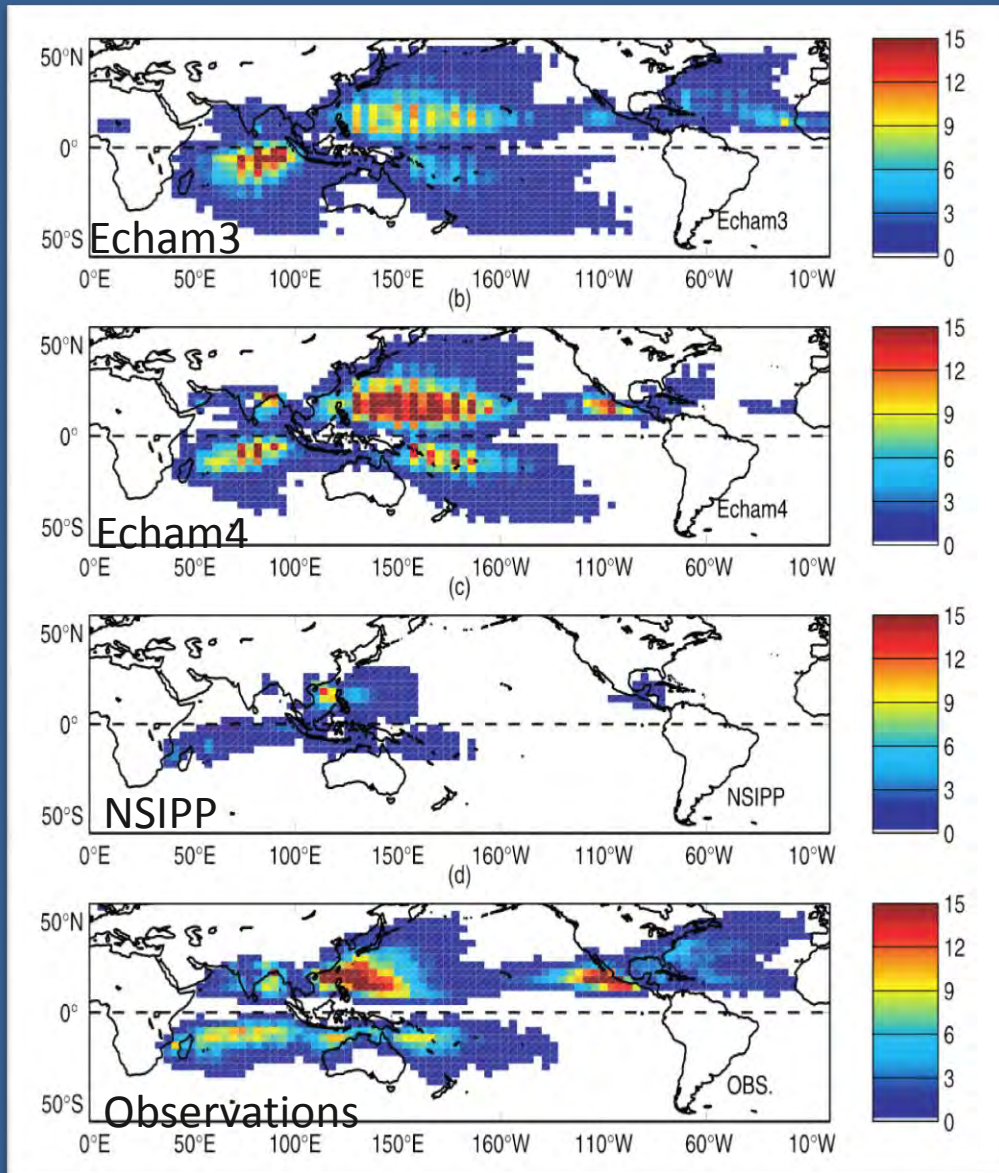


FIG. 4. Ensemble mean number of storms per year detected using basin-independent thresholds ( $V$ ), the basin-dependent threshold (3), and observations (\*) for (a) the Atlantic and (b) the Australia.

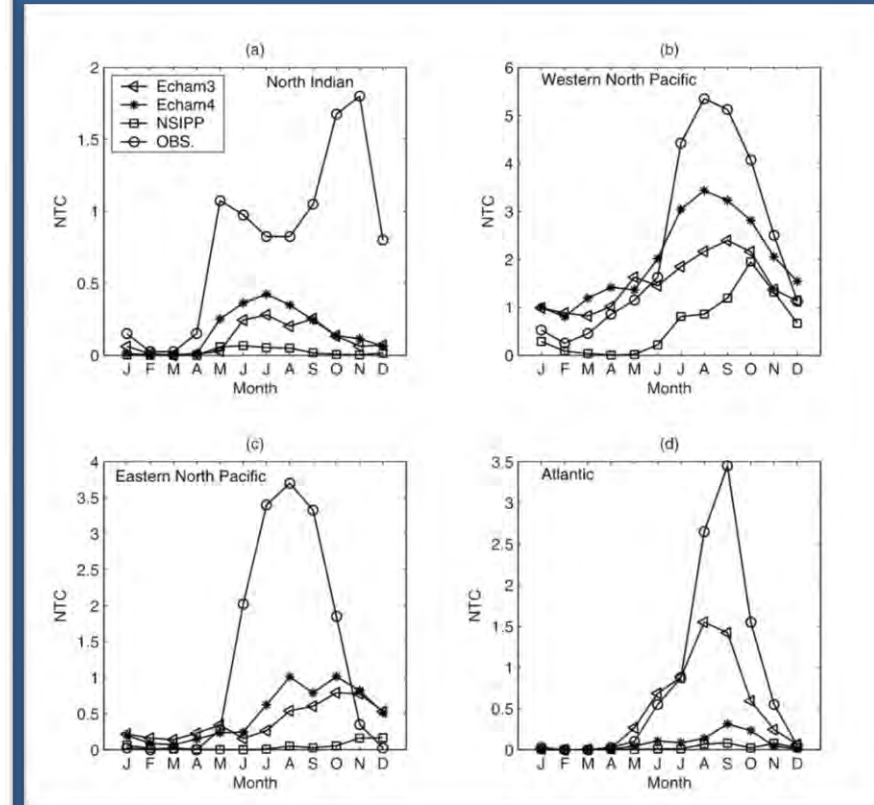
# TC dynamical seasonal forecasts



# More recent GCM studies: 2000s



Track density



Number of tropical cyclones

Camargo, Barnston & Zebiak, Tellus, 2005

# North Atlantic TC variability in current high-res GCMs used for climate system research: Atmosphere-only

1 JANUARY 2013

STRACHAN ET AL.

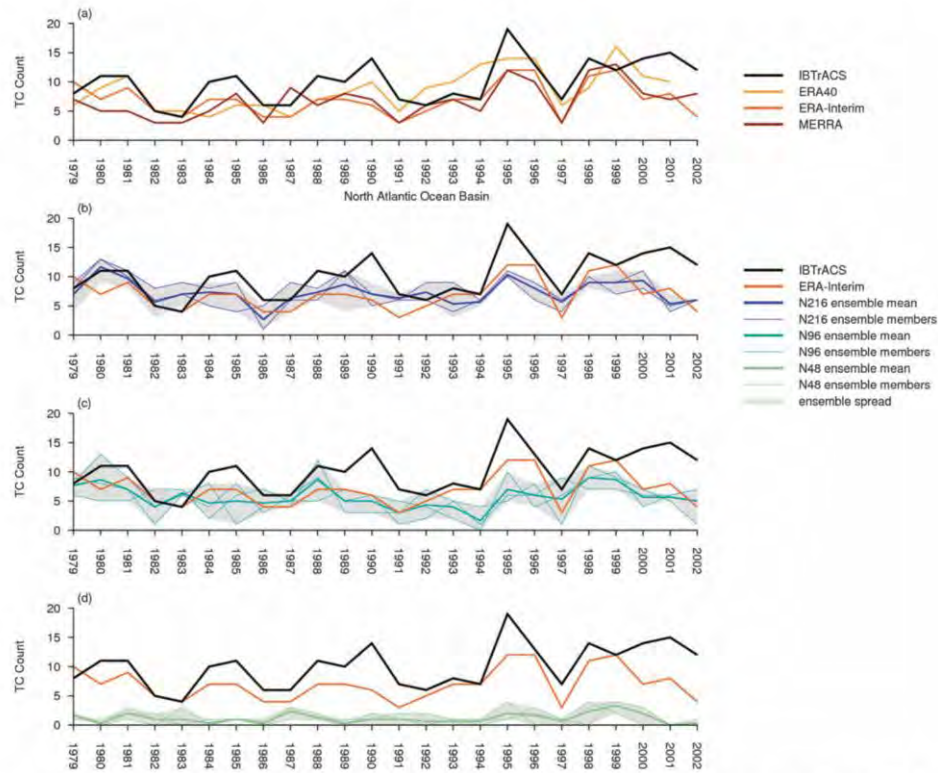


FIG. 7. Tropical cyclone interannual variability for North Atlantic basin for (a) reanalysis products; (b) the N216 GCM ensemble; (c) the N96 GCM ensemble; and (d) the N48 GCM ensemble. For (b)–(d) thick lines show the GCM ensemble mean, thin lines show the individual ensemble members, and the gray envelope shows the ensemble spread.

GFDL GCM:  
Zhao et al.  
J. Clim. (2009)

MBER 2009

ZHAO ET AL.

North Atlantic

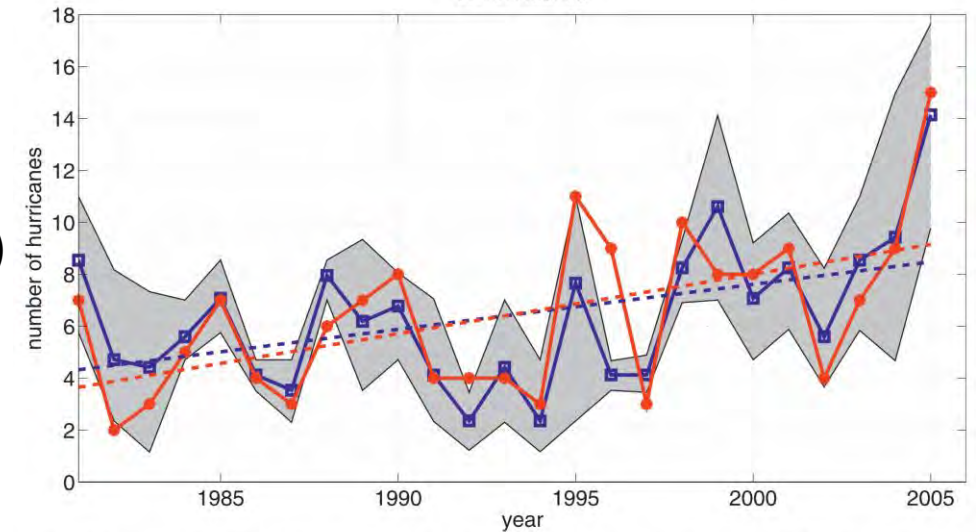
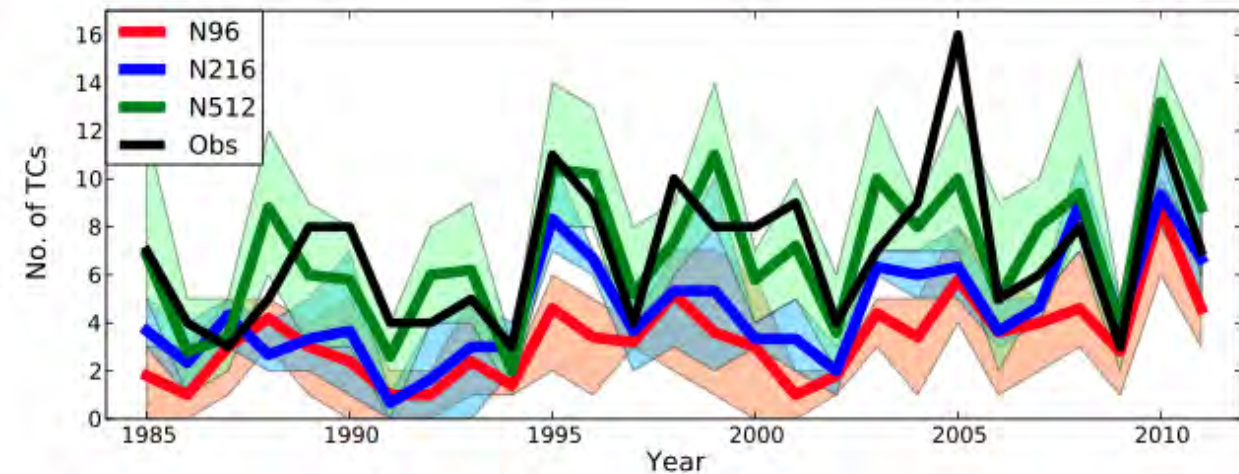


FIG. 7. Interannual variation of hurricane numbers for North Atlantic from 1981 to 2005. IBTrACS observations (Kruk et al. 2010) (red) and four-member ensemble mean (blue); shaded area shows the simulated maximum and minimum number for each year from the four-member integrations. Model time series are normalized by time-independent multiplicative factors so as to reproduce the observed total number. Dotted lines show observed and model (ensemble mean) linear trends.



HadGEM3: Roberts et al. J. Clim. (2015)

HadGEM1: Strachan et al. J. Clim. (2011-2013)

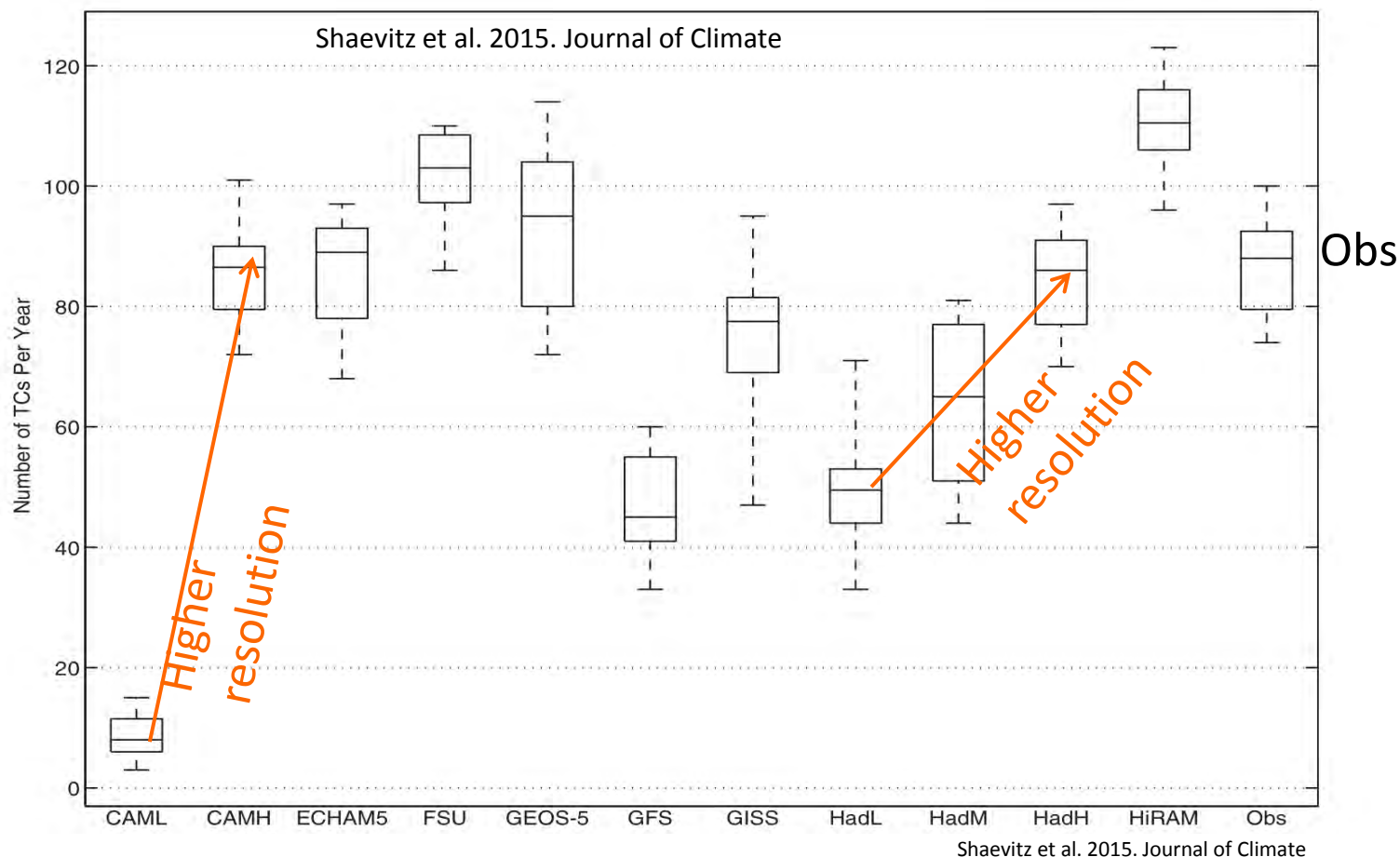
# Continuing issues

- Mean biases in TC climatology:
  - Number, intensity, tracks, size
- Mainly attributed to:
  - Model low-resolution
  - Mean biases in model environment
- Dependency on model characteristics:
  - Model physics: convection scheme
  - Dynamical core
- “*Hurricane-type vortices*”: if it walks like a duck and quacks like a duck...

# Tropical Cyclones “emerge” at high resolution

Results finally confirmed by the US CLIVAR Hurricane Working Group (HWG), via a **systematic** multi-model intercomparison:

- TC tracks and interannual variability in frequency are credibly represented at 20km;
- however, intensity is still underestimated by some of the GCMs at this resolution
- HRCM played a strong role in the first HWG; even stronger role in next phase



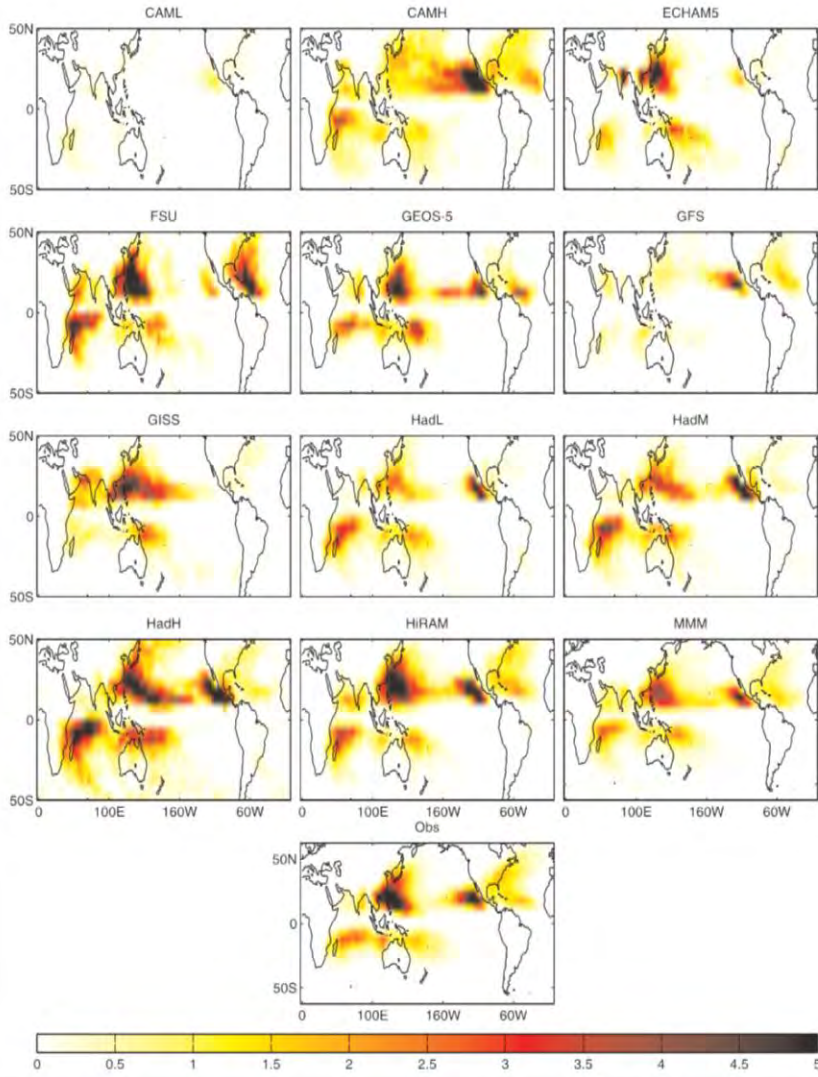
Distribution of the number of TCs per year

RESEARCH ARTICLE  
10.1002/2014MS000372

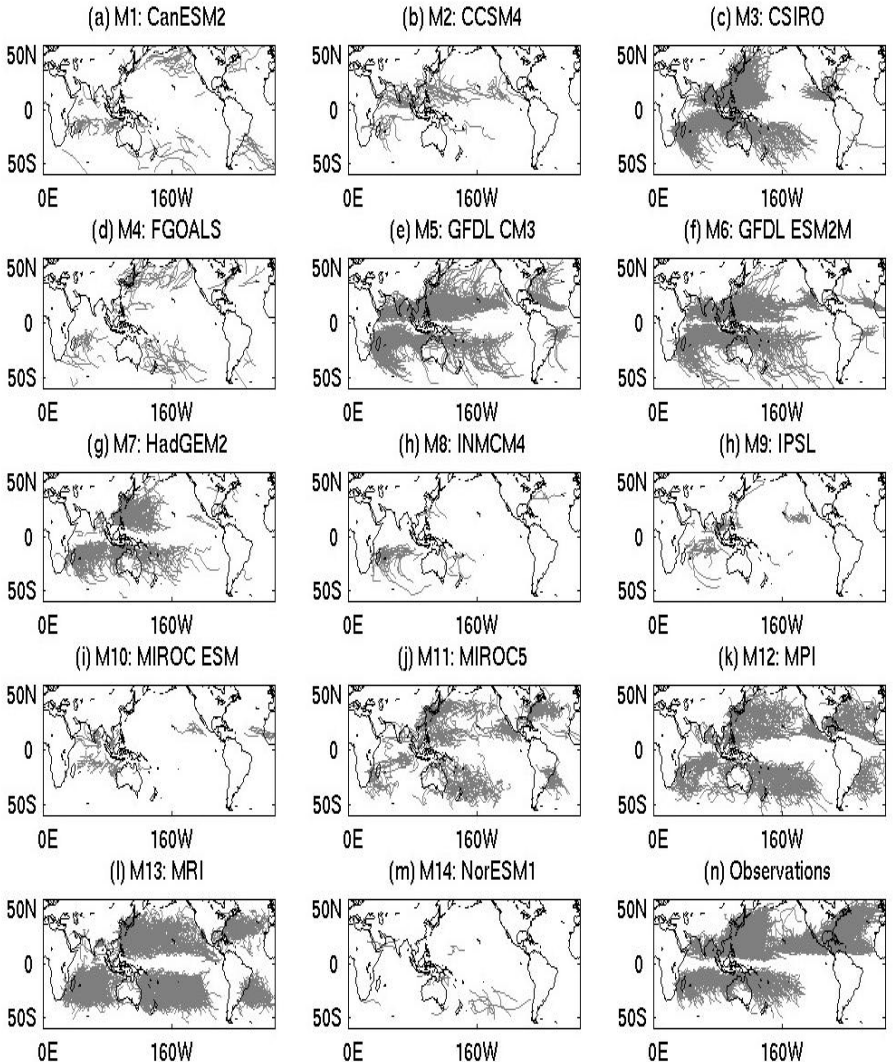
Characteristics of tropical cyclones in high-resolution models in the present climate

**Key Points:**  
• Multimodel comparison of tropical cyclone activity in global climate models  
• Geographic distribution of the TC

Daniel A. Shaevitz<sup>1</sup>, Suzana J. Camargo<sup>2</sup>, Adam H. Sobel<sup>1,2,3</sup>, Jeffrey A. Jonas<sup>4,5</sup>, Daehyun Kim<sup>2,6</sup>, Arun Kumar<sup>7</sup>, Timothy E. LaRow<sup>8</sup>, Young-Kwon Lim<sup>9,10</sup>, Hiroyuki Murakami<sup>11</sup>, Kevin A. Reed<sup>12</sup>, Malcolm J. Roberts<sup>13</sup>, Enrico Scoccimarro<sup>14,15</sup>, Pier Luigi Vidale<sup>16</sup>, Hui Wang<sup>7</sup>, Michael F. Wehner<sup>17</sup>, Ming Zhao<sup>18</sup>, and Naomi Henderson<sup>2</sup>



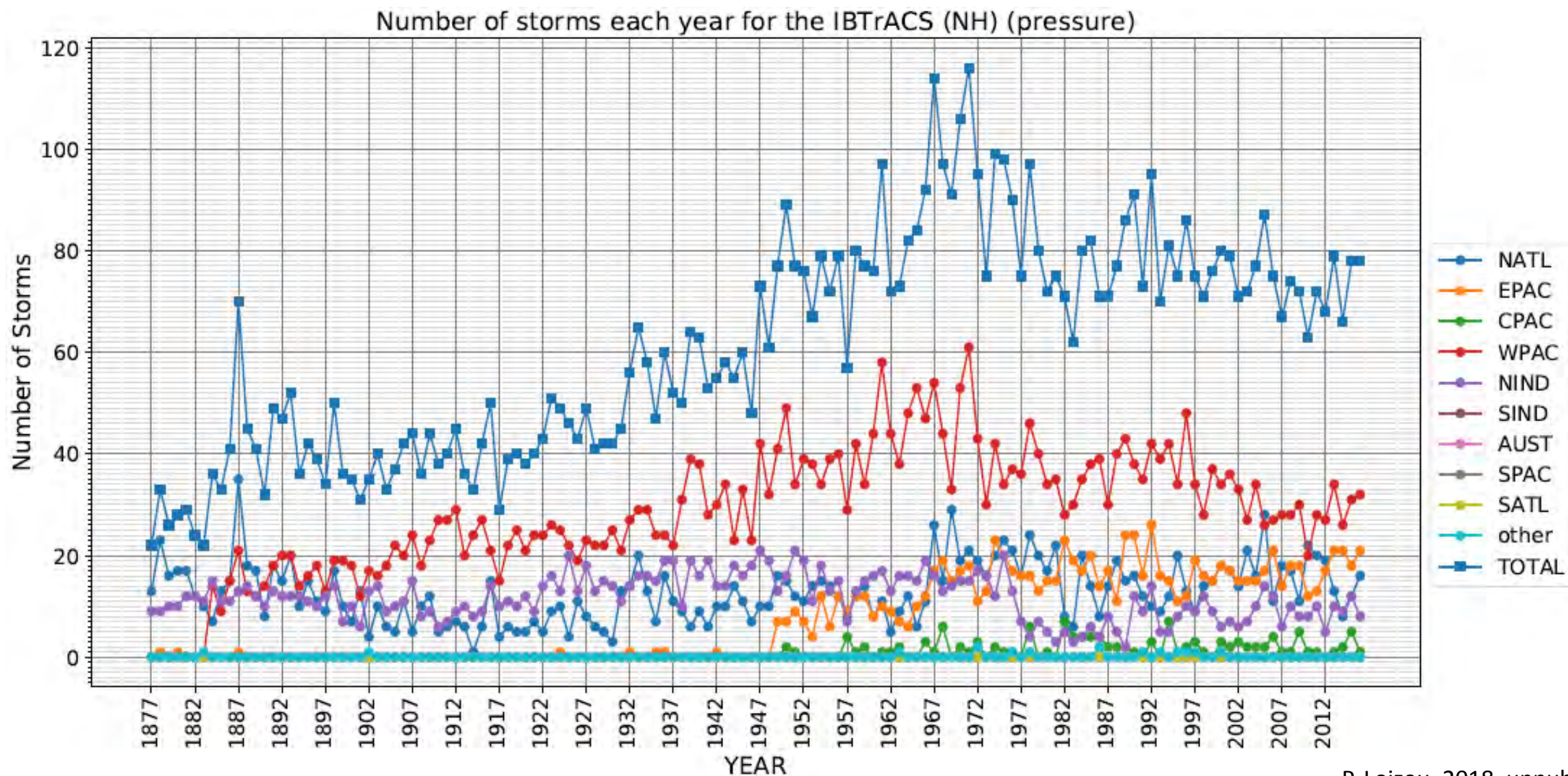
Track density



Tracks

What are the observed historical trends in key storm parameters and how well do the climate models resolve these?

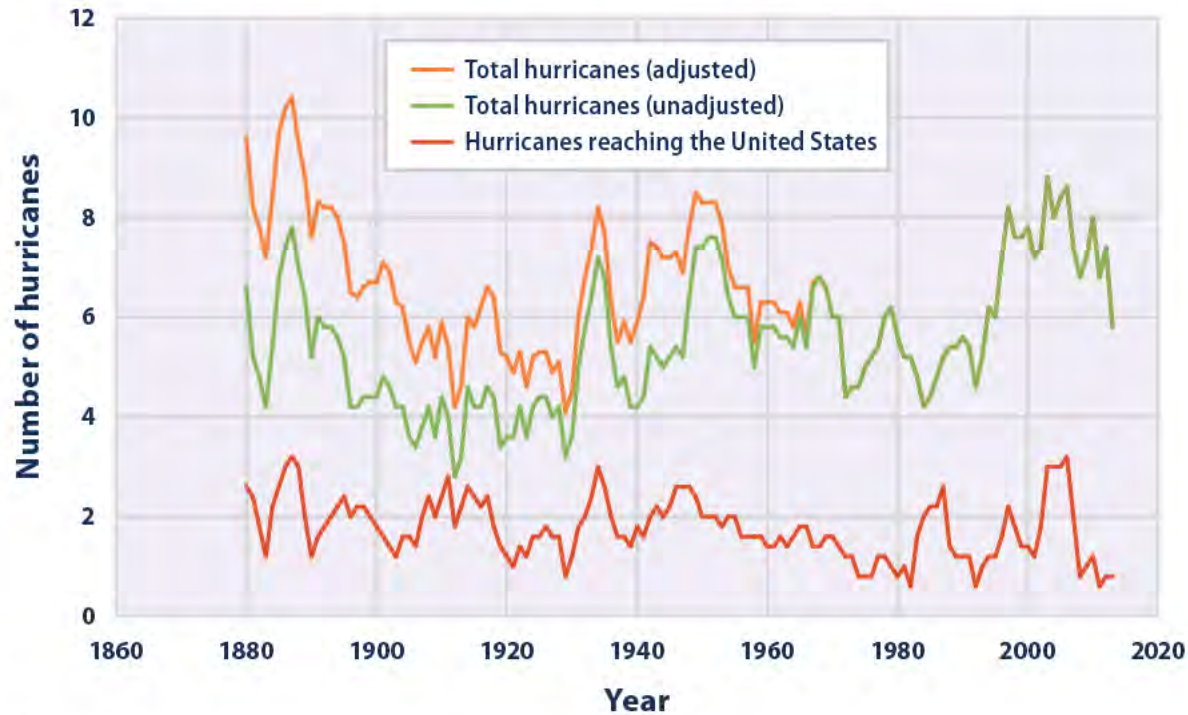
# TC observations are just a mess... (K. Hodges, 2018)



# There is a constant effort to re-visit and complete TC observations

Careful with fitting linear trends...

Number of Hurricanes in the North Atlantic, 1878–2015



Data sources:

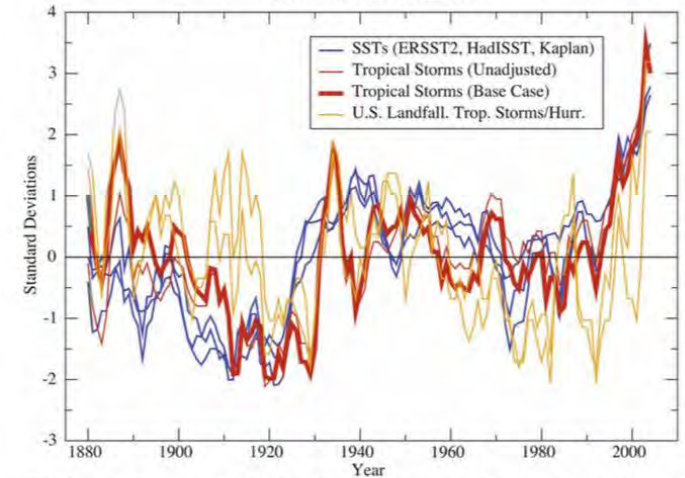
- NOAA (National Oceanic and Atmospheric Administration). 2016. The Atlantic Hurricane Database Re-analysis Project. [www.aoml.noaa.gov/hrd/hurdat/comparison\\_table.html](http://www.aoml.noaa.gov/hrd/hurdat/comparison_table.html).
- Vecchi, G.A., and T.R. Knutson. 2011. Estimating annual numbers of Atlantic hurricanes missing from the HURDAT database (1878–1965) using ship track density. J. Climate 24(6):1736–1746. [www.gfdl.noaa.gov/bibliography/related\\_files/gav\\_2010JCLI3810.pdf](http://www.gfdl.noaa.gov/bibliography/related_files/gav_2010JCLI3810.pdf).

For more information, visit U.S. EPA's "Climate Change Indicators in the United States" at [www.epa.gov/climate-indicators](http://www.epa.gov/climate-indicators).

Vecchi and Knutson 2011

Analyses of long time series (basically NATL) show that **substantial decadal variability is present** and needs to be considered in risk estimates and management actions.

(a) Tropical Atlantic SSTs, Tropical Storms, and U.S. Landfall Series  
5-yr running mean; Normalized



(b) Tropical Atlantic SSTs, Trop. Storms and Landfall Series: Detrended  
5-yr running mean; Normalized

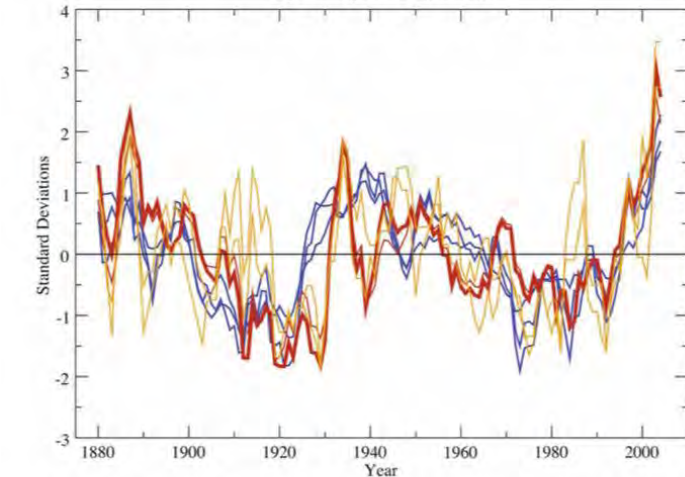


FIG. 10. (a) The 5-yr running mean normalized Atlantic MDR SST indices from three different reconstructions of SST (see text) overlaid on Atlantic TC counts. Blue curves are the three MDR SST reconstructions. The heavy red curve is the base-case TC count. The light red curve is the unadjusted TC count. The orange curves are U.S. landfalling tropical storm and hurricane count series from HURDAT. Curves in (b) have been detrended using ordinary least squares best fits.

Vecchi and Knutson 2008

DATASET	PERIOD	MODEL RESOLUTION	DATA GRID
IBTrACS	1877-2017		
ERA-Interim	1979-2016	TL255L60 (80 km)	512 × 256
MERRA	1979-2015	1/2° × 2/3° L72 (55 km)	540 × 361
MERRA-2	1980-2016	Cubed sphere (50 km)	576 × 361
NCEP	1979-2016	T382L64 (38 km)	720 × 361
JRA-25	1979-2013	T106L40 (120 km)	288 × 145
JRA-55	1958-2017	TL319L60 (55 km)	288 × 145

# TCs in 6 re-analyses

Main finding: re-analyses are able to credibly reproduce TCs in the higher categories, but are challenged in the TD, TS and CAT1 categories.

15 JULY 2017

HODGES ET AL.

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TABLE 3. Storms that match and do not match with IBTrACS in the NH by storm category, for each reanalysis, storms identified by the objective detection method applied to the reanalysis tracks and, in parentheses, the direct matching method, performed in [section 3a](#). Values are number per year.

Category		ERA-Interim	JRA-25	JRA-55	NCEP-CFSR	MERRA	MERRA-2
TD	Match	2.91 (7.94)	3.26 (7.94)	5.24 (8.03)	3.50 (8.00)	2.29 (7.85)	3.48 (7.67)
	No match	5.56 (0.53)	5.21 (0.53)	3.24 (0.44)	4.97 (0.47)	6.18 (0.62)	4.91 (0.73)
TS	Match	11.85 (22.38)	18.62 (22.53)	18.32 (22.53)	14.76 (22.32)	9.85 (22.44)	14.24 (22.45)
	No match	11.85 (1.32)	5.09 (1.18)	5.38 (1.18)	8.94 (1.38)	13.85 (1.26)	9.73 (1.52)
CAT1	Match	8.74 (12.23)	11.18 (12.23)	11.17 (12.24)	10.09 (12.12)	7.74 (12.21)	9.76 (12.33)
	No match	3.44 (0.00)	1.00 (0.00)	1.00 (0.00)	2.09 (0.06)	4.44 (0.00)	2.55 (0.00)
CAT2	Match	5.29 (6.35)	6.15 (6.38)	6.00 (6.35)	5.82 (6.38)	4.76 (6.35)	5.64 (6.39)
	No match	1.06 (0.00)	0.21 (0.00)	0.35 (0.00)	0.53 (0.00)	1.59 (0.00)	0.73 (0.00)
CAT3	Match	6.15 (7.00)	6.91 (7.06)	6.82 (7.03)	6.71 (7.06)	5.82 (7.03)	6.42 (7.06)
	No match	0.88 (0.03)	0.12 (0.00)	0.21 (0.00)	0.32 (0.00)	1.21 (0.00)	0.64 (0.00)
CAT4	Match	5.97 (6.79)	6.76 (6.79)	6.71 (6.74)	6.47 (6.79)	5.76 (6.76)	6.48 (6.76)
	No match	0.82 (0.00)	0.03 (0.00)	0.09 (0.06)	0.32 (0.00)	1.03 (0.03)	0.33 (0.06)
CAT5	Match	1.09 (1.12)	1.12 (1.12)	1.12 (1.12)	1.12 (1.12)	1.03 (1.12)	1.09 (1.09)
	No match	0.03 (0.00)	0.00 (0.00)	0.00 (0.00)	0.00 (0.00)	0.09 (0.00)	0.00 (0.00)

We created a complementary TC database that can be used to assist risk assessment, by:

- increasing the sample size and
- providing physically based estimates of model uncertainty.

# There is no trace of TC trends in the re-analyses: there is, however, substantial and interesting variability

NA NORTH ATLANTIC (P)

WP WEST PACIFIC (P)

IBTrACS

ERA-Interim

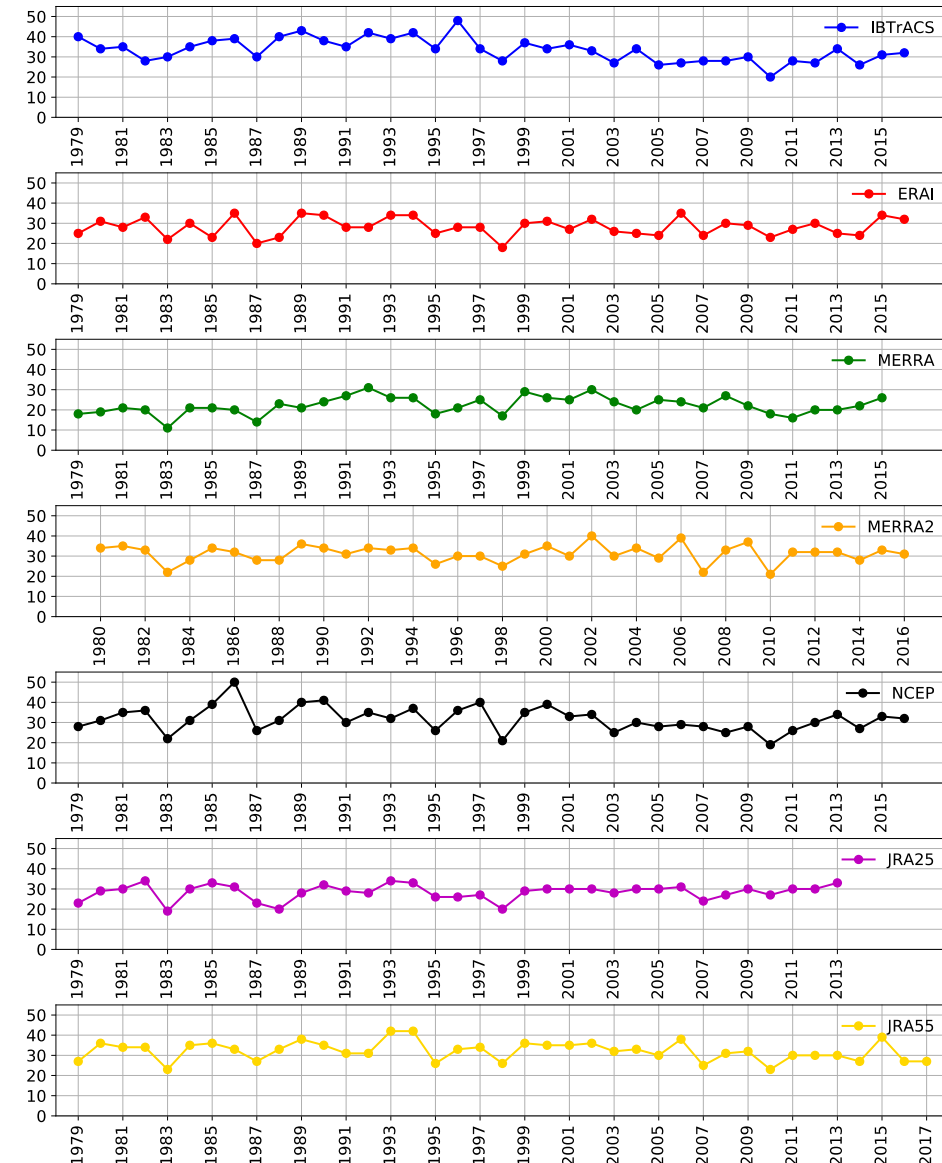
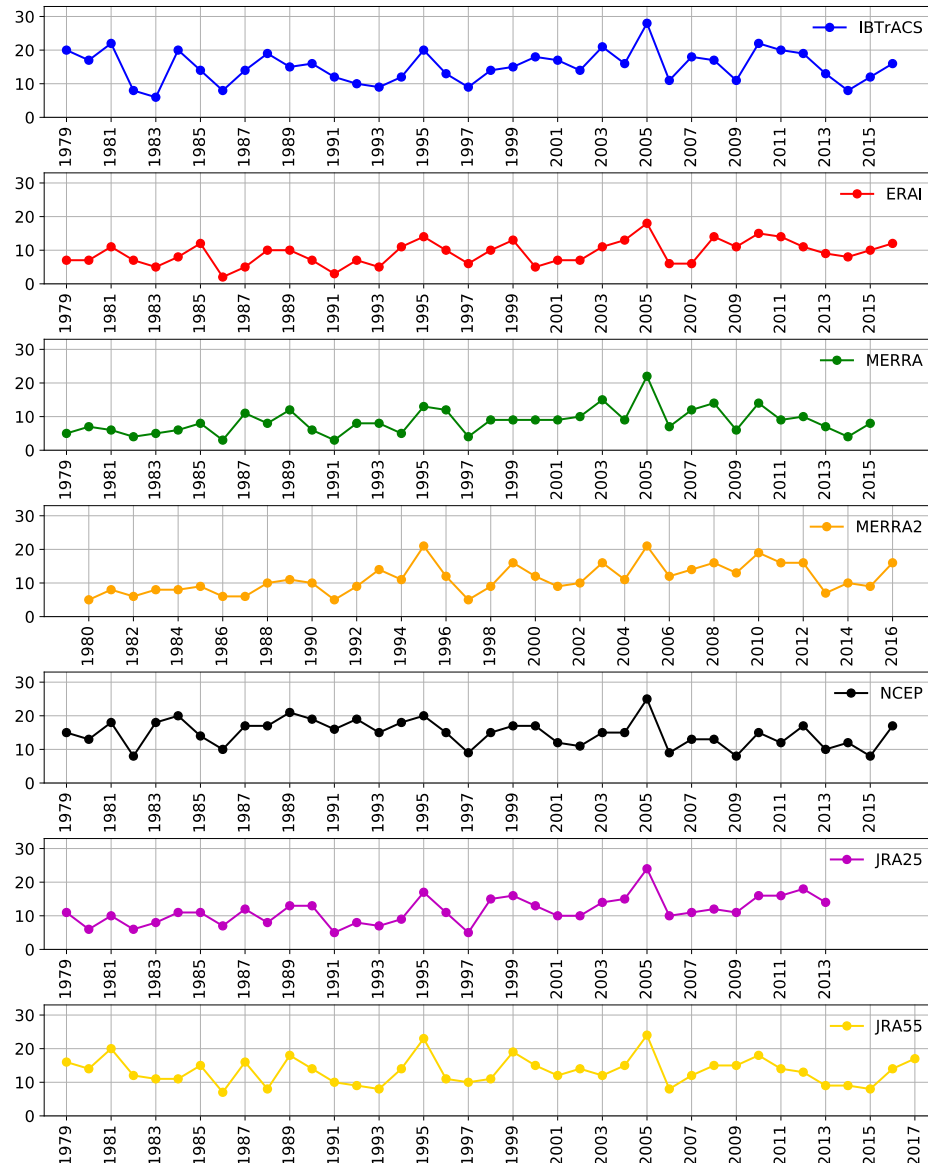
MERRA

MERRA2

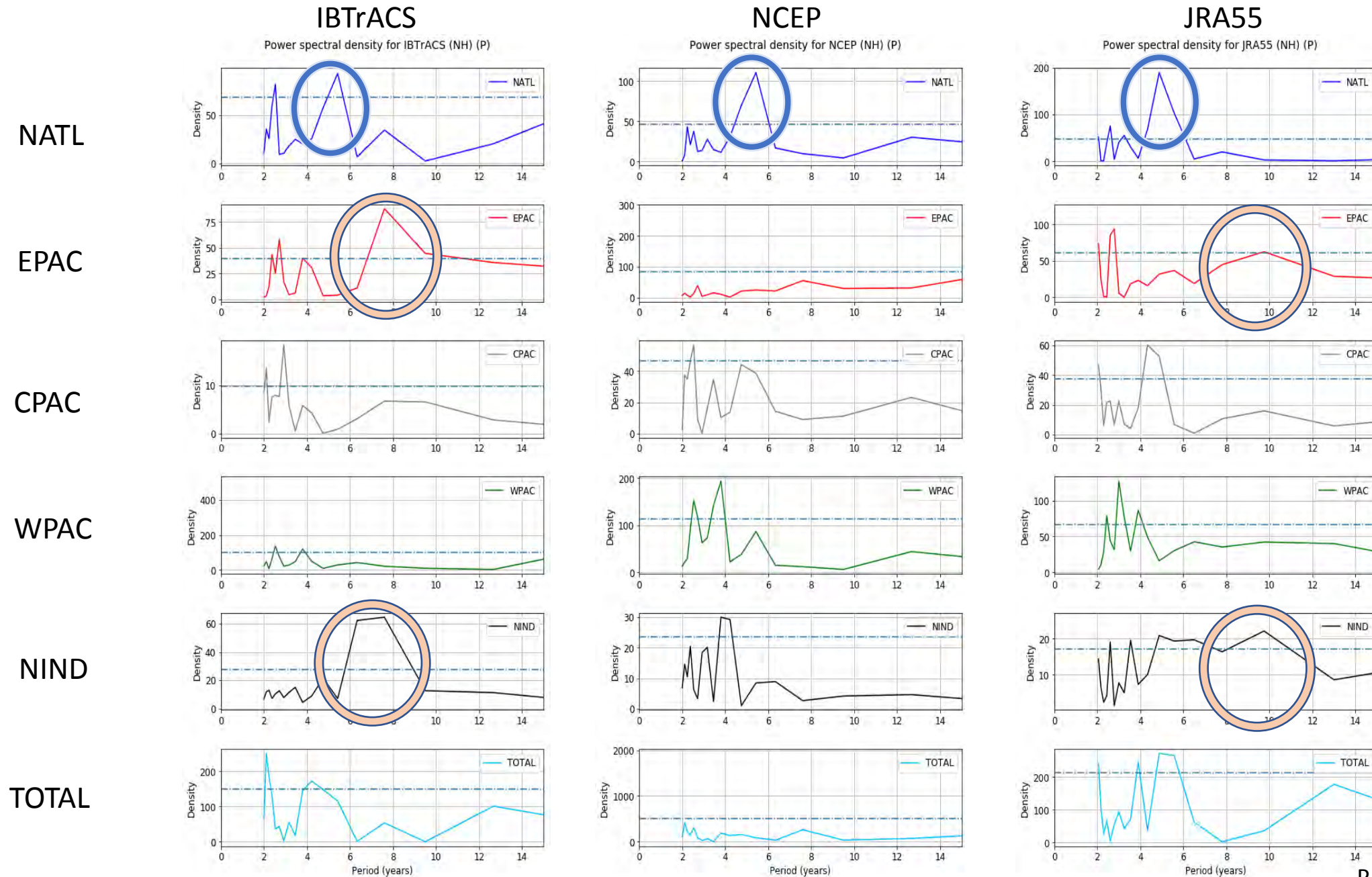
NCEP

JRA25

JRA55

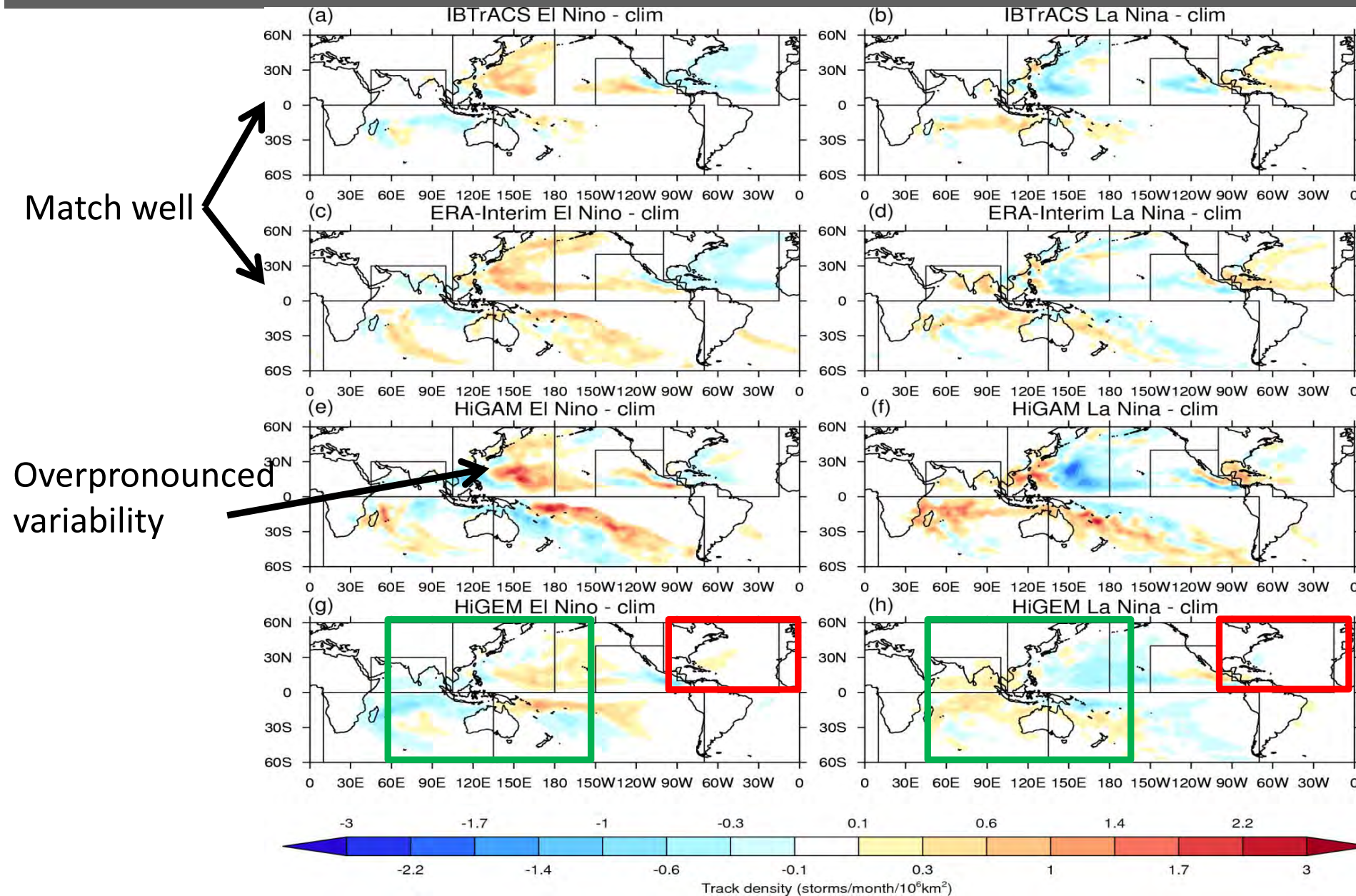


# Power spectra of TC time series in different basins, from 3 data sets

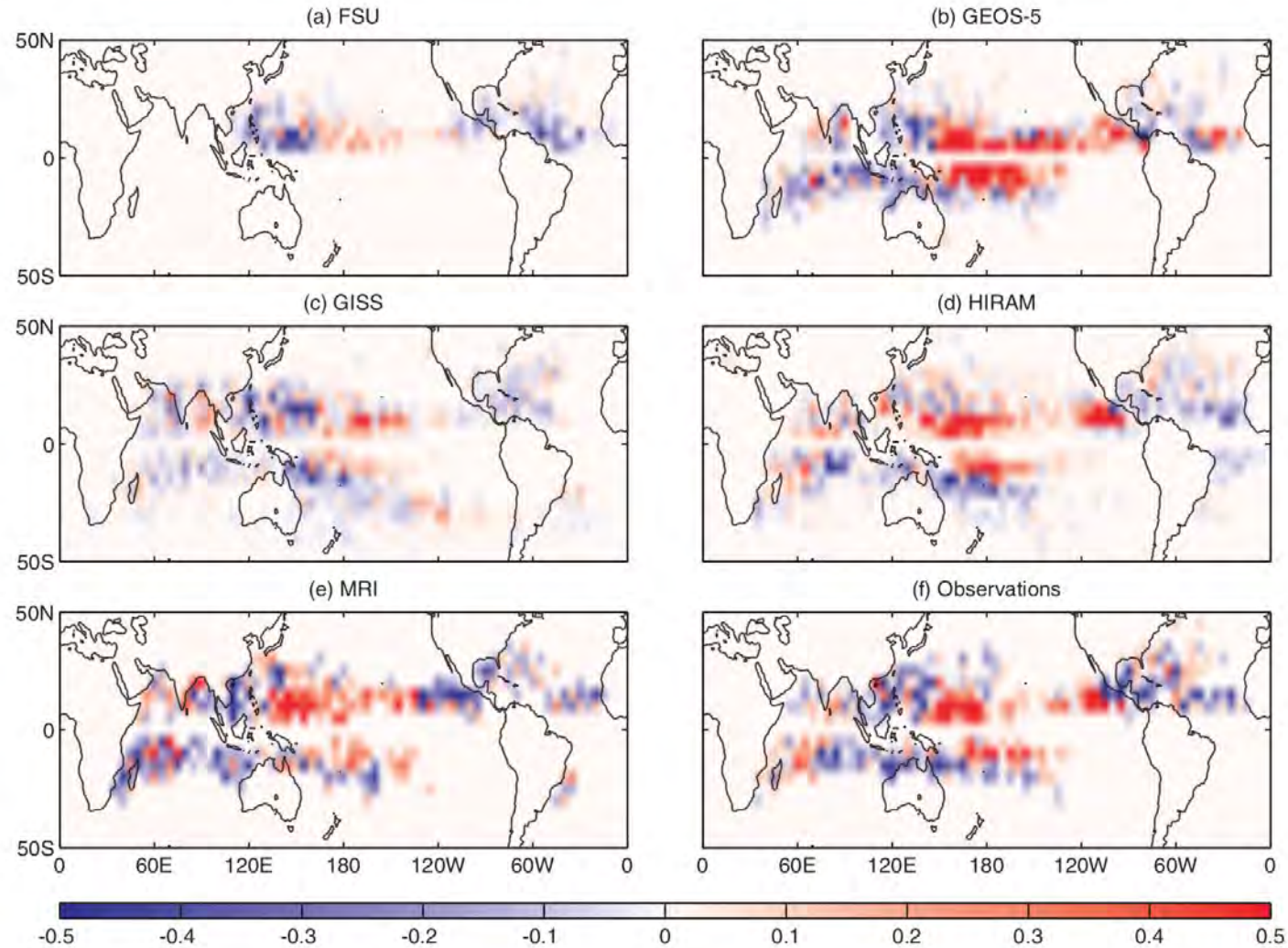


*Will we see changes to the inter-dependency between La Niña / El Niño cycles and tropical storm intensity or frequency?*

# ENSO-TC: track density anomalies

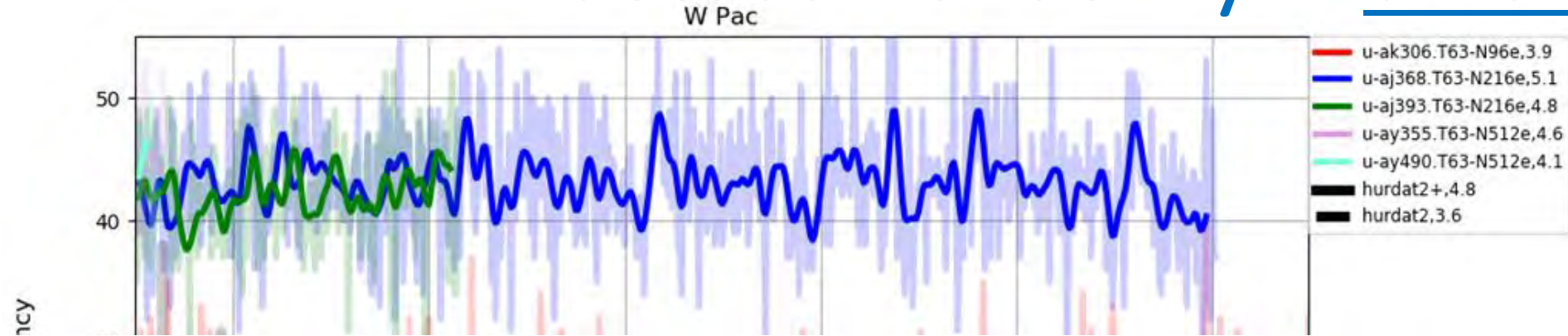


# TCs and ENSO in CLIVAR HWG exercise



**Figure 11.** Difference of TC genesis density in El Niño and La Niña in models and observations. The genesis density is defined as the mean TC formation per  $5^\circ \times 5^\circ$  box per year.

# PRIMAVERA: decadal variability in unforced runs



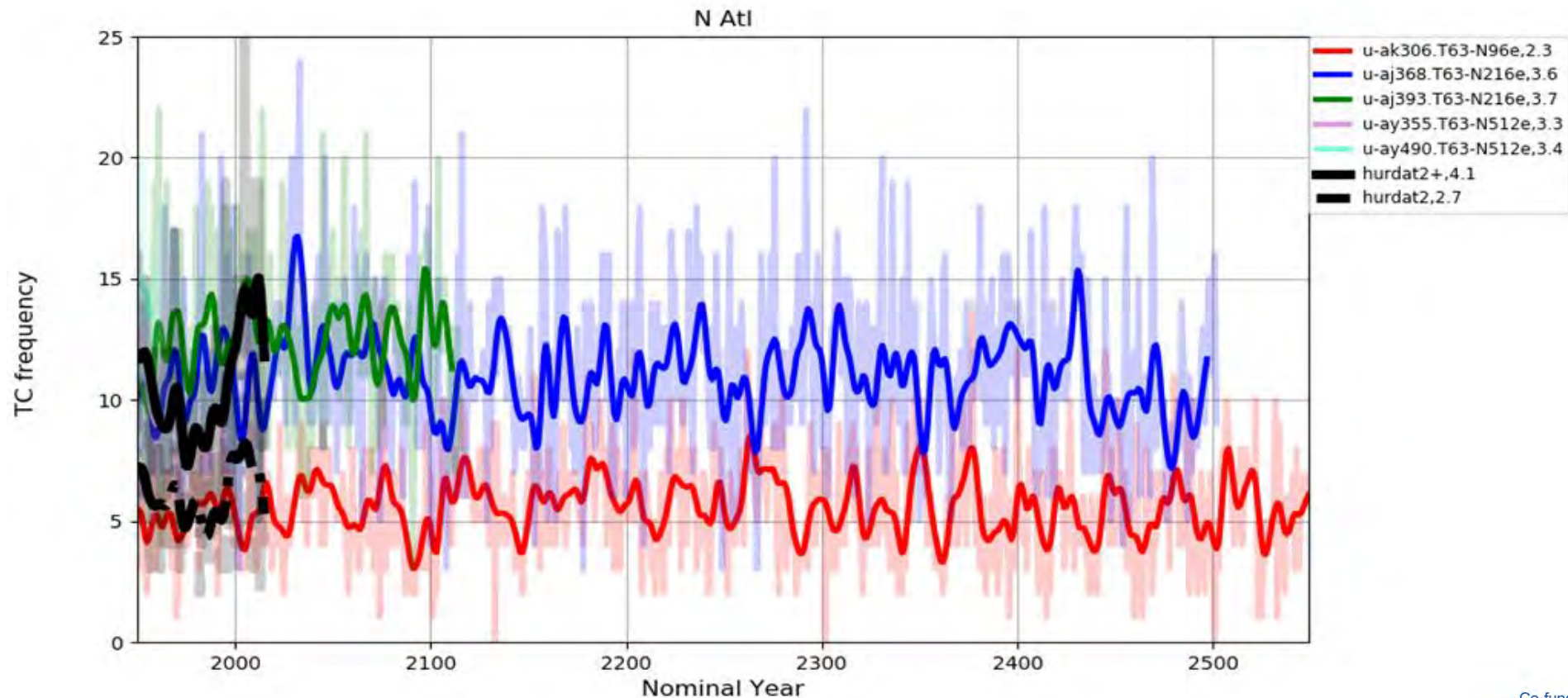
N96-ORCA1

N216-ORCA025

N216-ORCA12

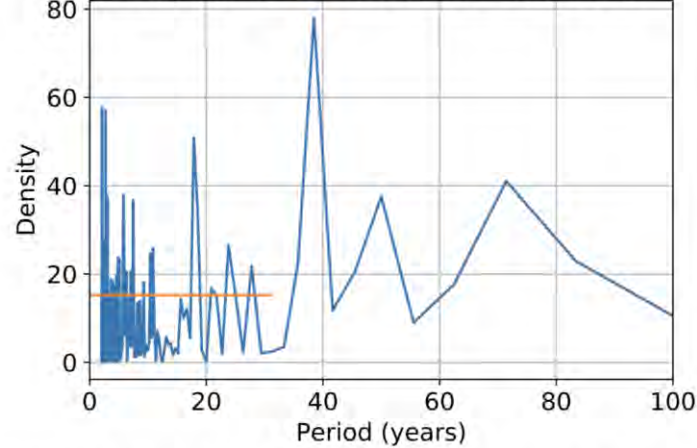
N512-ORCA025

N512-ORCA12

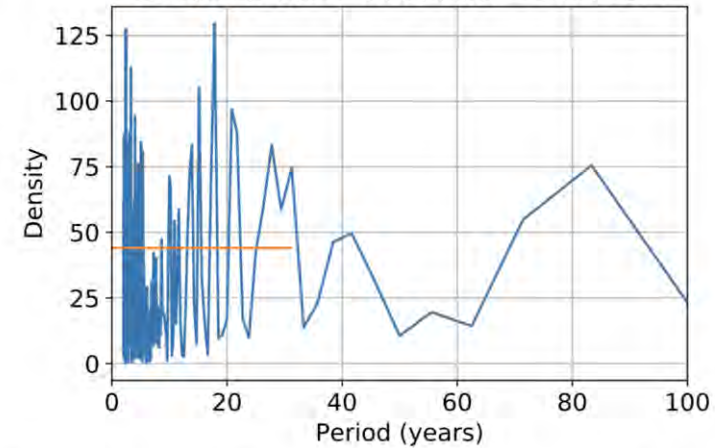


# Power spectra of TC time series in different basins, from HadGEM3 simulations

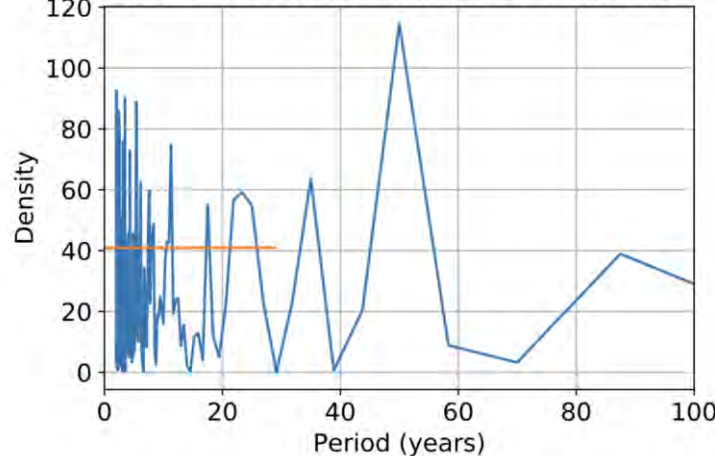
Power Spectral Density for TC Counts - u-ak306 NATL  
GC3.1-LL N96 - ORCA025 - COUPLED



Power Spectral Density for TC Counts - u-ak306 WPAC  
GC3.1-LL N96 - ORCA025 - COUPLED



Power Spectral Density for TC Counts - u-aj368 NATL  
GC3.1-MM N216 - ORCA025 - COUPLED



Power Spectral Density for TC Counts - u-aj368 WPAC  
GC3.1-MM N216 - ORCA025 - COUPLED

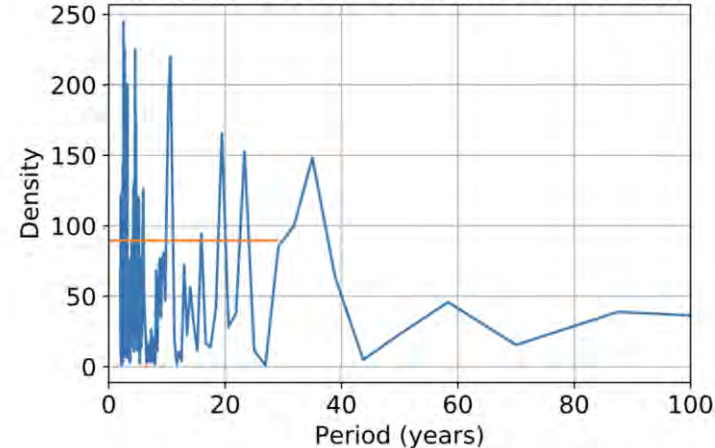


Figure 5: Results of the power spectrum analysis for the coupled model experiments. The two upper figures show the results for the N96- experiment for the NATL (on the left) and WPAC (on the right) regions. The two lower figures show the results for the N216-experiments for the same regions.

What are the projected changes to track paths, intensity, structure, frequency and potential seasonal extensions?

Is there regional variation in these climate model projections?

# What will happen to TCs in the future?

Typhoons will migrate poleward ... and a NA hurricane reduction

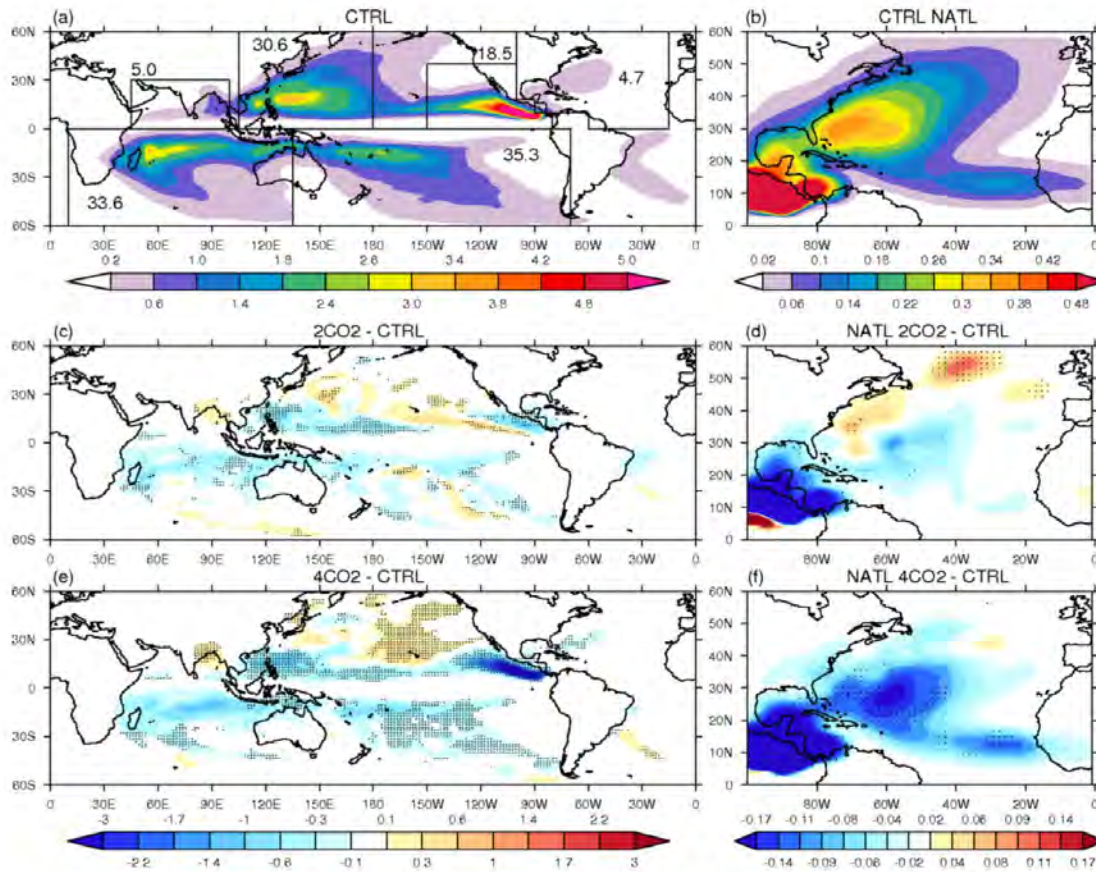
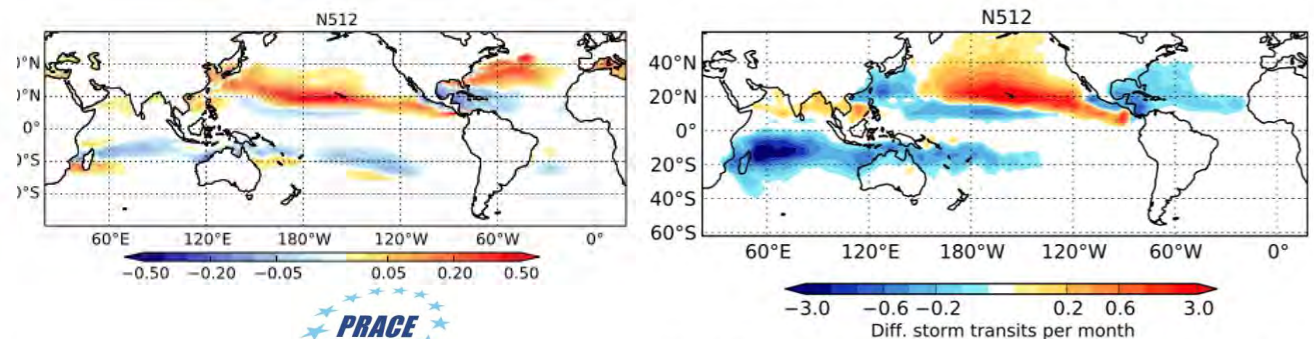


FIG. 2. Tropical cyclone track density, same as figure 1, for (a) HiGEM present-day simulation (b) The same as for (a) but North Atlantic (c) 2CO<sub>2</sub> - present-day simulation (d) North Atlantic 2CO<sub>2</sub> - present-day simulation (e) 4CO<sub>2</sub> - present-day simulation and (f) North Atlantic 4CO<sub>2</sub> - present-day simulation. Stippling shows where changes are outside the range of 5x30-year present-day simulations.

Bell et al. J. Clim. 2012, idealised HiGEM simulations

GPI-based estimates agree in the Pacific, albeit not in the Atlantic



## 2012 UPSCALE MODELLING CAMPAIGN

JOURNAL OF CLIMATE

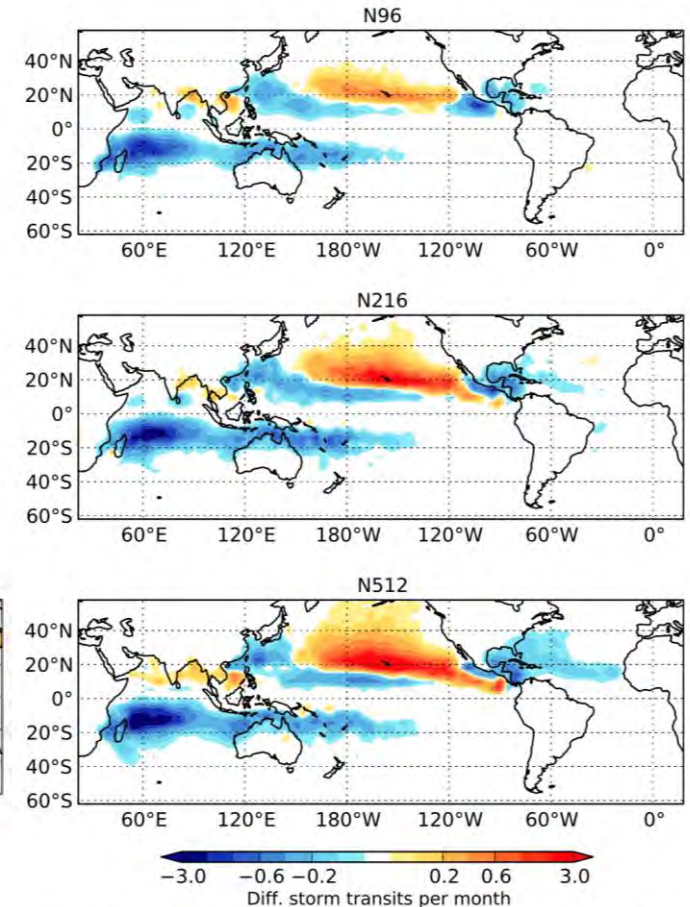
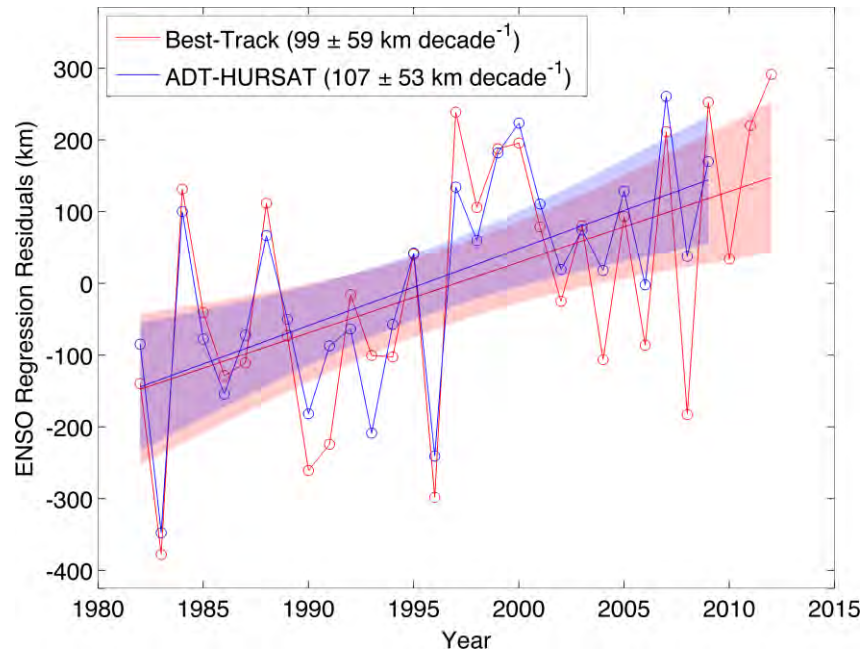


FIG. 12. Change in tropical cyclone track density (storm transits per month per unit area equivalent to a 4° spherical cap) between the future climate and present climate integrations for the whole 1986-2010 period and for the whole ensemble at each model resolution: (top) N96, (middle) N216, and (bottom) N512.

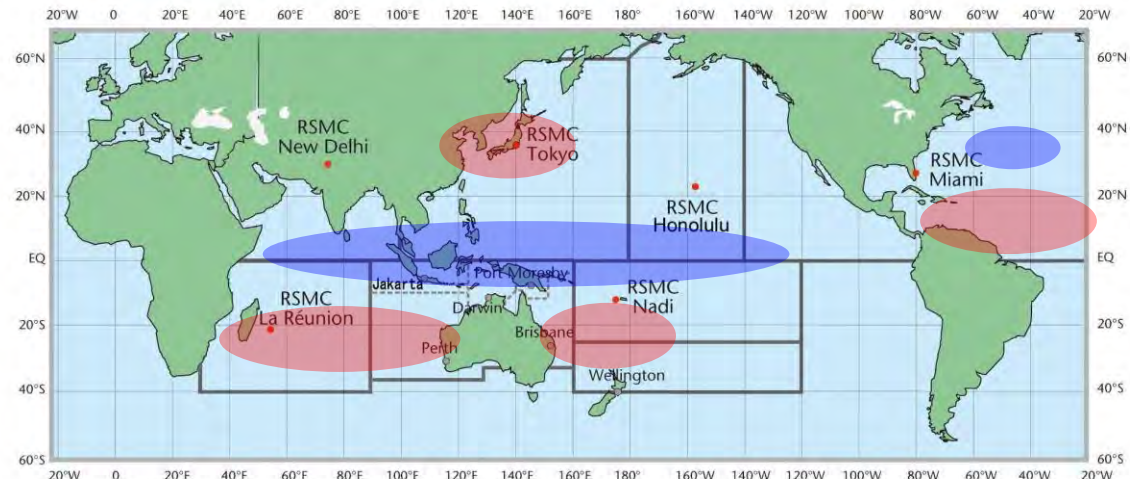
Roberts et al. 2015. Journal of Climate, RCP 8.5 scenario

## Back to observations: annual-mean latitude of peak TC intensity

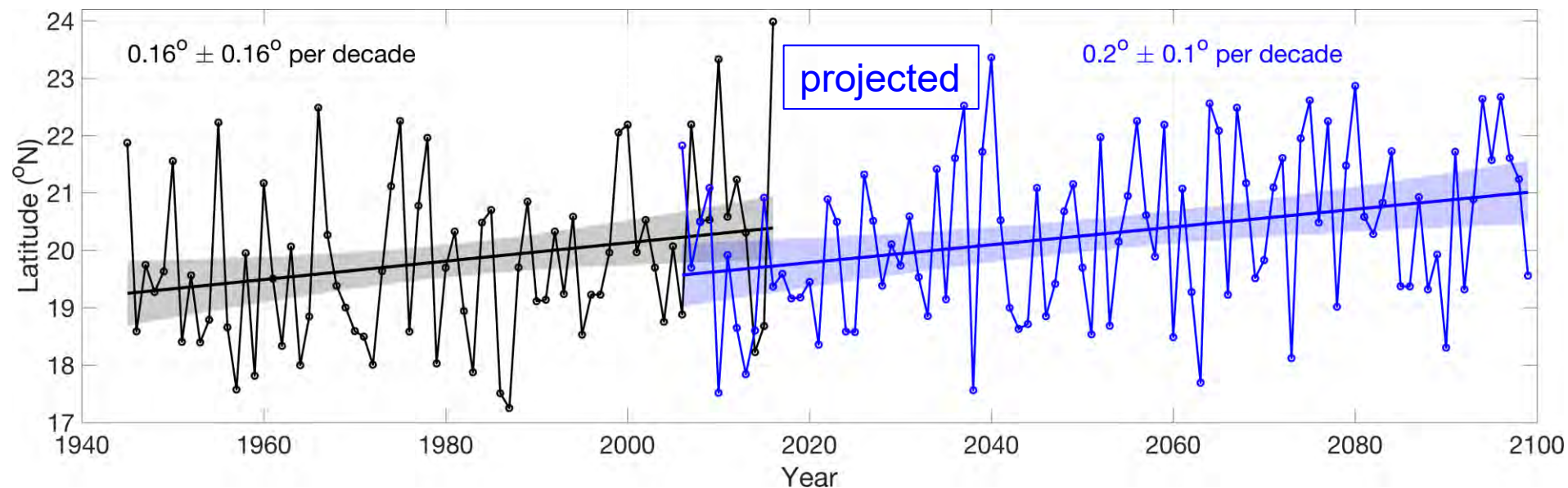


The migration rate is consistent with the independently-measured rate of tropical expansion, which has been partly attributed to human activity.

Kossin, J. P., K. A. Emanuel, and G. A. Vecchi, 2014: The poleward migration of the location of tropical cyclone maximum intensity. *Nature*, **509**, 349-352.

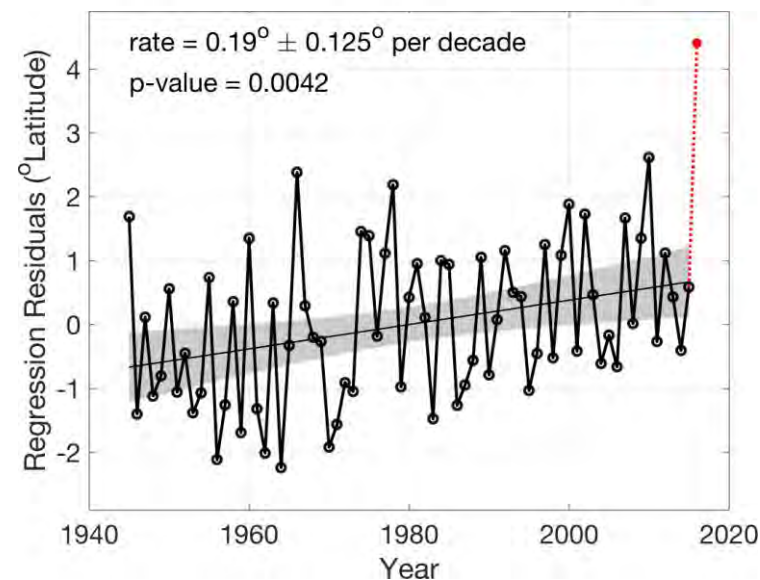


# Longer-term observed trends and projections

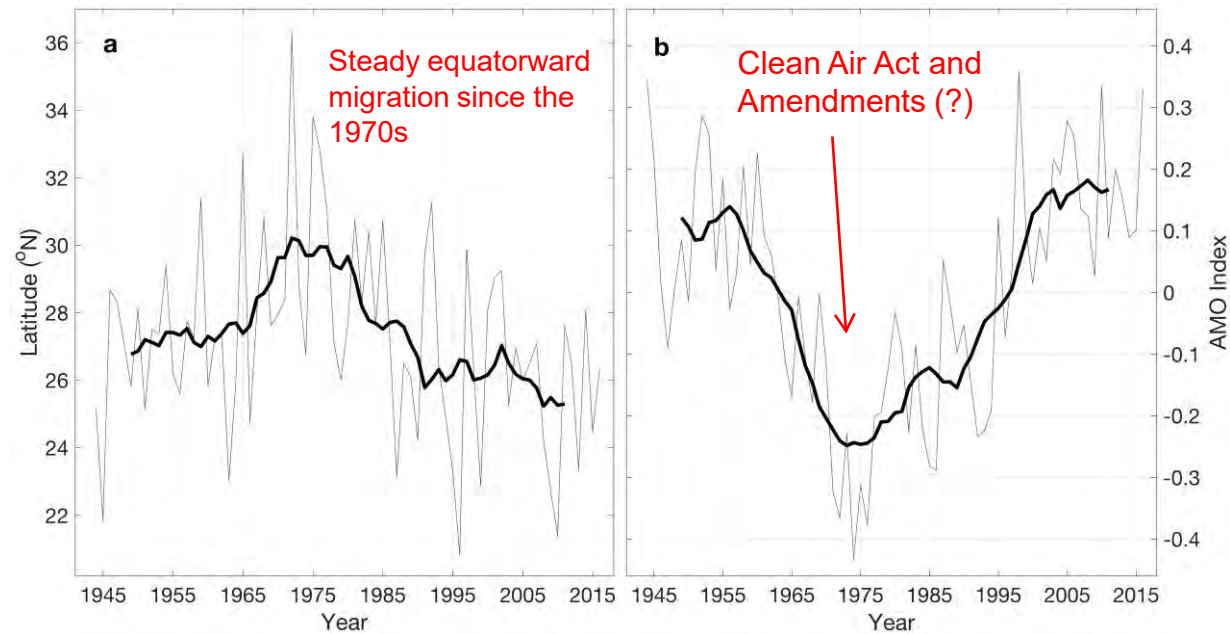


Can we separate part of the trend from natural variability?

Western North Pacific known dominant modes of natural variability:  
ENSO (inter-annual)  
PDO (decadal)

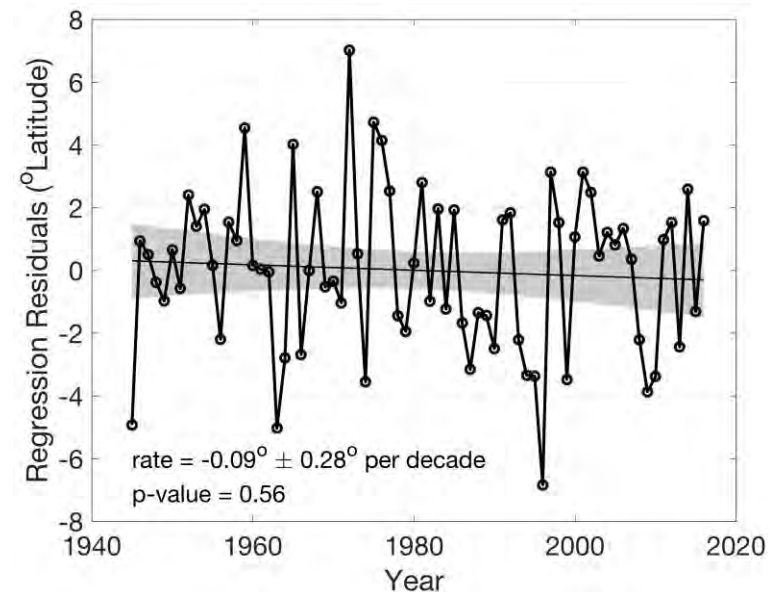


# Back to observations: North Atlantic migration of TCs



**No long-term trend.**

But there is uncertainty about what's driving the Atlantic Multi-decadal Oscillation



# What will happen to TCs in the future?

## Models agree on an overall reduction.

What is still controversial:

- a) regional distribution
- b) changes by category

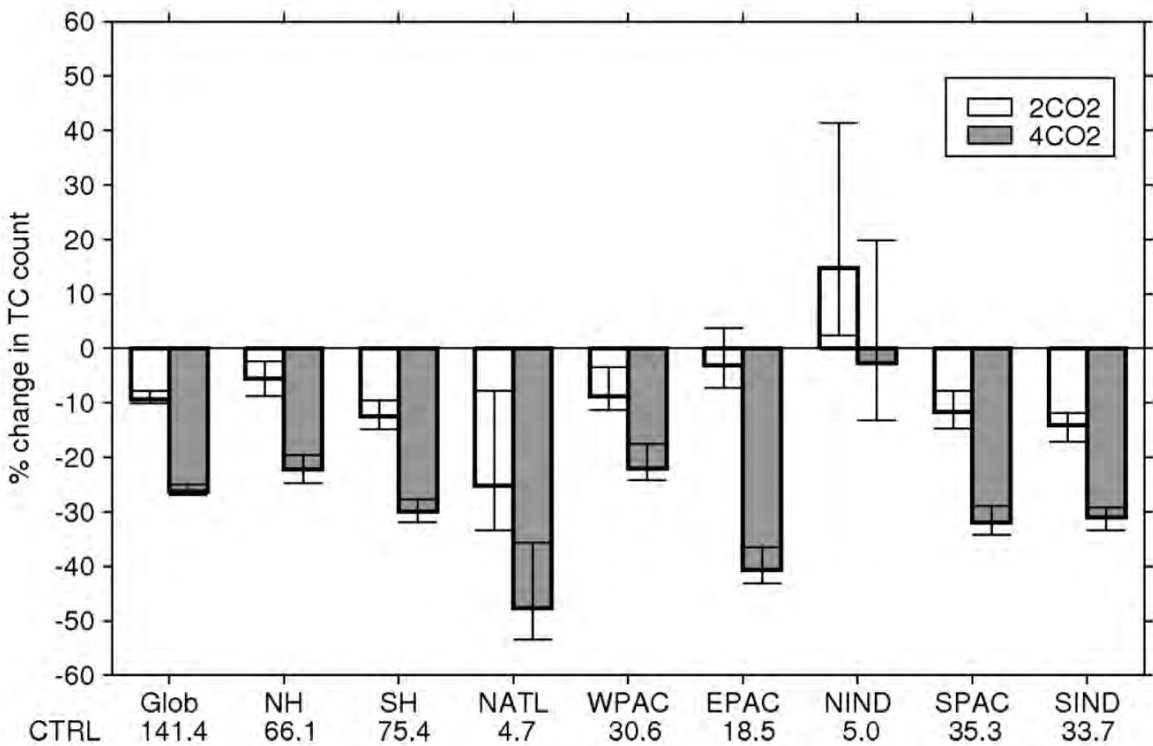


FIG. 3. Percentage change of annual tropical cyclone counts. The error bars denote the maximum and minimum 5x30-year present-day simulations. The present-day climatology is shown at the bottom.

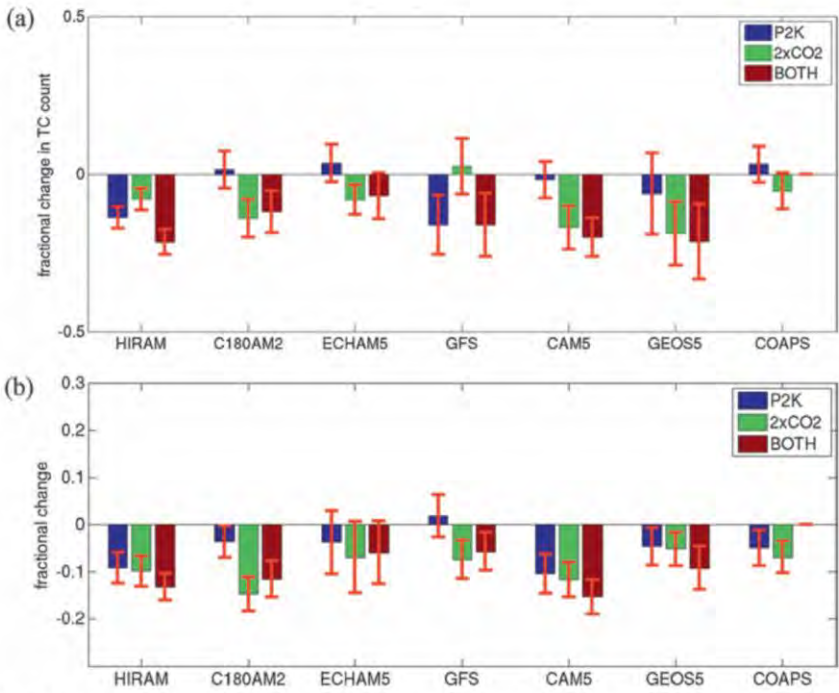


FIG. 3. Comparison between changes in (a) TC formation for various models for the 2K (P2K) and 2CO2 experiments vs (b) TC genesis as weighted by changes in midtropospheric vertical velocity, as described in the text. (From Zhao et al. 2013a.)

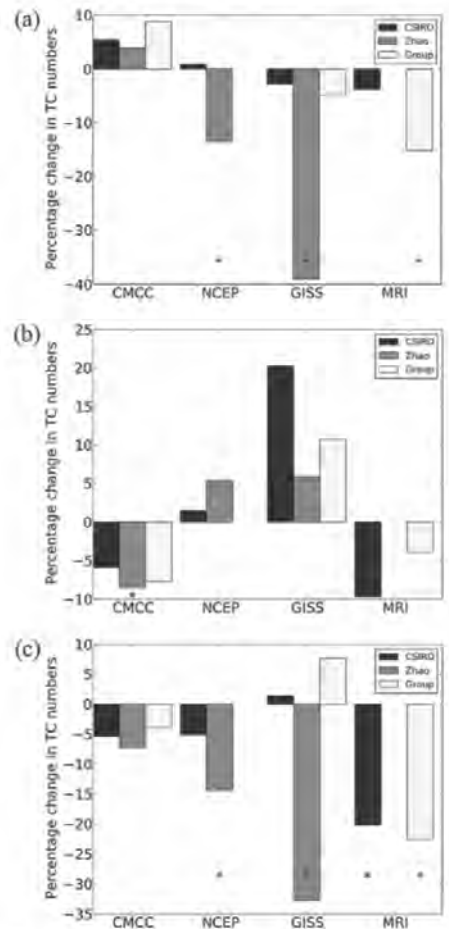


FIG. 6. Percentage change in TC numbers in each model for the three altered climate experiments relative to the present-day experiment, as tracked by the Commonwealth Scientific and Industrial Research Organisation (CSIRO), Zhao, and individual group tracking schemes, after homogenization in (a) duration, (b) wind speed, and (c) latitude of formation. Asterisks indicate statistical significance to at least the  $p = 0.05$  level.

# Projected changes in intensity due to anthropogenic climate change: still a controversial issue

2015 KNUTSON ET AL.

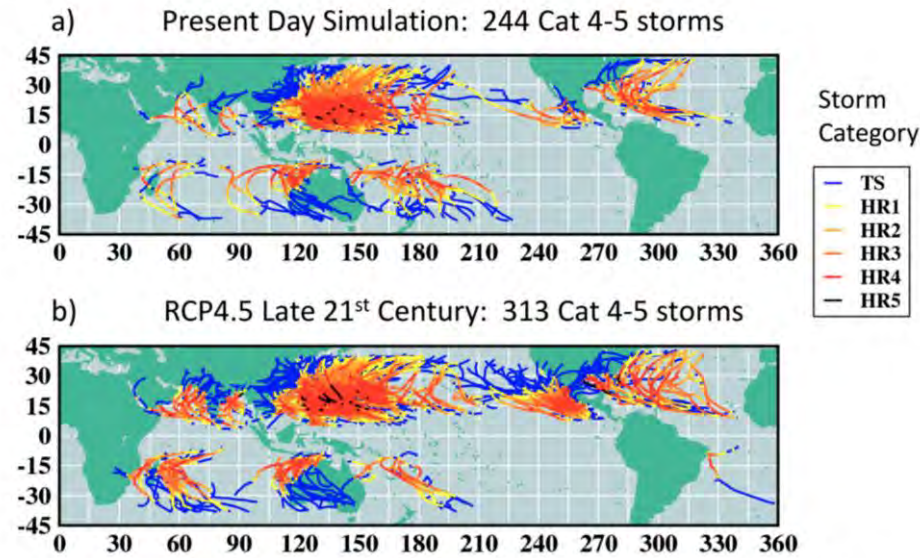


FIG. 7. Tracks of simulated cat 4–5 tropical cyclones for (a) present-day or (b) late-twenty-first-century (RCP4.5; CMIP5 multimodel ensemble) conditions. Simulated tropical cyclone tracks were obtained using the GFDL hurricane model to resimulate (at higher resolution) the tropical cyclone cases originally obtained from the HiRAM C180 global mode. Storm categories or intensities are shown over the lifetime of each storm, according to the Saffir–Simpson scale. The categories are depicted by the track colors, varying from tropical storm (blue) to category 5 (black; see legend).

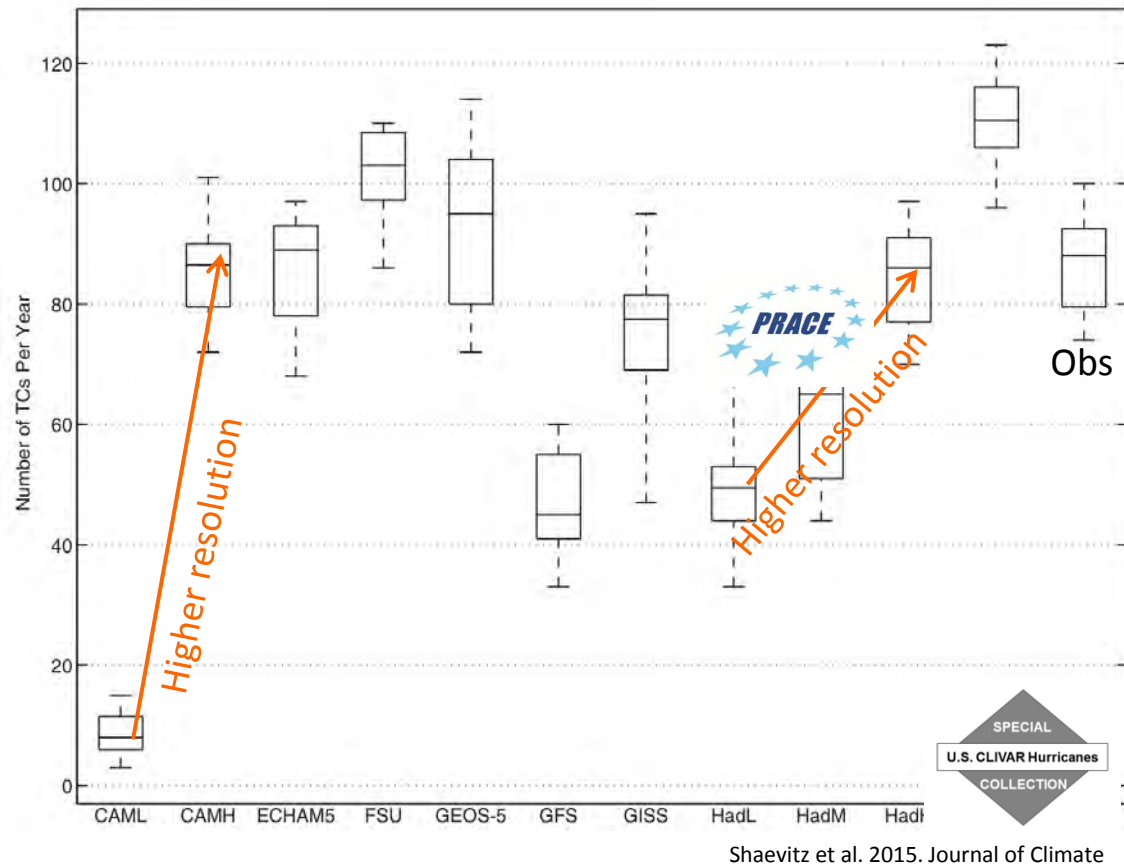
TABLE 3. Tropical cyclone activity (percent change) statistics from downscaling experiments for CMIP5 multimodel ensembles (future vs present day). The future scenarios use RCP4.5 averaged conditions for late twenty-first century and are compared to the “present-day” simulations for 1982–2005 climatological SST conditions. “Cat” refers to Saffir–Simpson intensity category (1–5) with “cat 0” signifying less than hurricane strength. Rain rate is the average rain rate within 100 km of the storm center, including all tropical cyclones (not just 10% rainiest). PDI is power dissipation index in units of  $10^9 \text{ m}^3 \text{ s}^{-2}$ . “Hur (wind > 65)” refers to hurricanes with surface wind speeds greater than or equal to  $65 \text{ m s}^{-1}$ . “Maxwnd\_tc” and “maxwnd\_hur” are percent changes of mean lifetime-maximum intensities for all tropical cyclones (wind speed >  $17.5 \text{ m s}^{-1}$ ) or hurricanes (wind speed >  $33 \text{ m s}^{-1}$ ). The  $p$  values, for a null hypothesis of no change from present to future, are given in the line below each percent change entry. These use a two-sided Mann–Whitney–Wilcoxon test for all frequency or days-of-occurrence metrics, and a one-sided test (for increase) for the intensity and rain rate metrics. Bold values indicate significance at the  $p < 0.05$  level. “Inf” refers to cases where no occurrences were simulated in the present-day run while some were simulated in the future runs, indicating an infinite percent increase.

Variable	Global	North Atlantic	Northeast Pacific	Northwest Pacific	North Indian	South Indian	Southwest Pacific
No. of TC (cat 0–5)	<b>−16.4</b>	−9.4	<b>16.3</b>	<b>−34.5</b>	19.5	<b>−26.1</b>	<b>−36.6</b>
$p$ value	<b>&lt;0.01</b>	0.39	<b>0.02</b>	<b>&lt;0.01</b>	0.07	<b>&lt;0.01</b>	<b>&lt;0.01</b>
No. of hur (cat 1–5)	<b>−16.6</b>	−17.5	<b>19.3</b>	<b>−31.6</b>	<b>25.6</b>	<b>−28.4</b>	<b>−40.6</b>
$p$ value	<b>&lt;0.01</b>	0.16	<b>0.01</b>	<b>&lt;0.01</b>	<b>0.04</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>
No. of hur (cat 3–5)	1.8	2.7	<b>83.7</b>	<b>−16.9</b>	21.7	−8.3	<b>−50.6</b>
$p$ value		0.76	<b>&lt;0.01</b>	<b>&lt;0.01</b>	0.14	0.42	<b>&lt;0.01</b>
No. of hur (cat 4–5)	<b>28.3</b>	42.1	<b>337.5</b>	−6.5	200.0	63.6	<b>−58.3</b>
$p$ value	<b>&lt;0.01</b>	0.63	<b>&lt;0.01</b>	0.59	−0.05	0.07	<b>0.01</b>
No. of hur (wind > 65)	<b>59.3</b>	125.0	Inf	16.7	Inf	Inf	0.0
$p$ value	<b>0.01</b>	0.36		0.63			1.00
ACE	−15.1	−9.7	<b>44.2</b>	<b>−26.9</b>	23.2	<b>−28.8</b>	<b>−41.7</b>
$p$ value	0.15	0.29	<b>&lt;0.01</b>	<b>&lt;0.01</b>	0.13	<b>&lt;0.01</b>	<b>&lt;0.01</b>
PDI	−9.7	−3.1	<b>52.7</b>	<b>−22.7</b>	28.6	<b>−26.6</b>	<b>−43.9</b>
$p$ value	<b>&lt;0.01</b>	0.53	<b>&lt;0.01</b>	<b>&lt;0.01</b>	0.08	<b>&lt;0.01</b>	<b>&lt;0.01</b>
maxwnd_ts	<b>3.6</b>	0.4	<b>8.2</b>	<b>7.4</b>	3.4	1.8	−5.6
$p$ value	<b>&lt;0.01</b>	0.41	<b>&lt;0.01</b>	<b>&lt;0.01</b>	0.07	0.20	0.98
maxwnd_hur	<b>4.1</b>	<b>4.5</b>	<b>7.8</b>	<b>5.5</b>	1.6	<b>3.3</b>	−3.1
$p$ value	<b>&lt;0.01</b>	<b>0.04</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	0.23	<b>0.03</b>	0.88
Cat 4–5 days	<b>34.5</b>	175.4	<b>478.1</b>	9.7	<b>405.0</b>	55.4	<b>−52.5</b>
$p$ value	<b>&lt;0.01</b>	0.14	<b>&lt;0.01</b>	0.50	<b>0.04</b>	0.39	<b>0.03</b>
Rain rate_tc (cat 0–5)	<b>14.3</b>	<b>17.3</b>	<b>17.2</b>	<b>20.8</b>	<b>10.5</b>	<b>8.5</b>	−1.2
$p$ value	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>0.01</b>	<b>0.03</b>	0.66
Rain rate_hur (cat 1–5)	<b>13.4</b>	<b>20.5</b>	<b>14.4</b>	<b>15.5</b>	<b>12.8</b>	<b>11.1</b>	3.5
$p$ value	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	0.58
Rain rate_hur (cat 3–5)	<b>8.8</b>	<b>14.0</b>	<b>9.7</b>	<b>9.7</b>	<b>10.8</b>	<b>6.1</b>	10.6
$p$ value	<b>&lt;0.01</b>	<b>0.02</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	<b>&lt;0.01</b>	0.65
Rain rate_hur (cat 4–5)	<b>7.7</b>	9.4	<b>11.4</b>	<b>8.7</b>	21.4	−1.5	15.3
$p$ value	<b>&lt;0.01</b>	0.13	<b>&lt;0.01</b>	<b>&lt;0.01</b>	0.06	0.54	0.93
Delta SST (°C)		1.546	1.685	1.567	1.551	1.384	1.145

*What are the present model limitations and how are these being (or could be) resolved?*

# Tropical Cyclones “emerge” at high resolution

From US CLIVAR  
Hurricane Working Group (2015)



Our main question: **is this a robust result?**  
We need a multi-model, multi-resolution,  
ensemble approach

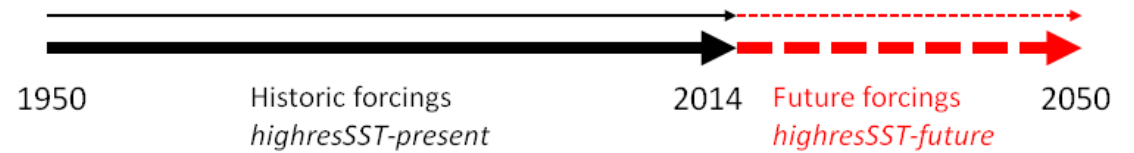
to

## CMIP6-HighResMIP TC simulations PRIMAVERA, 2018

Atmosphere-land-only, 1950-2014 (→ 2050)

Forced by observed SST and sea-ice and historic forcings (→ projected)

highresSST-present (→ highresSST-future)

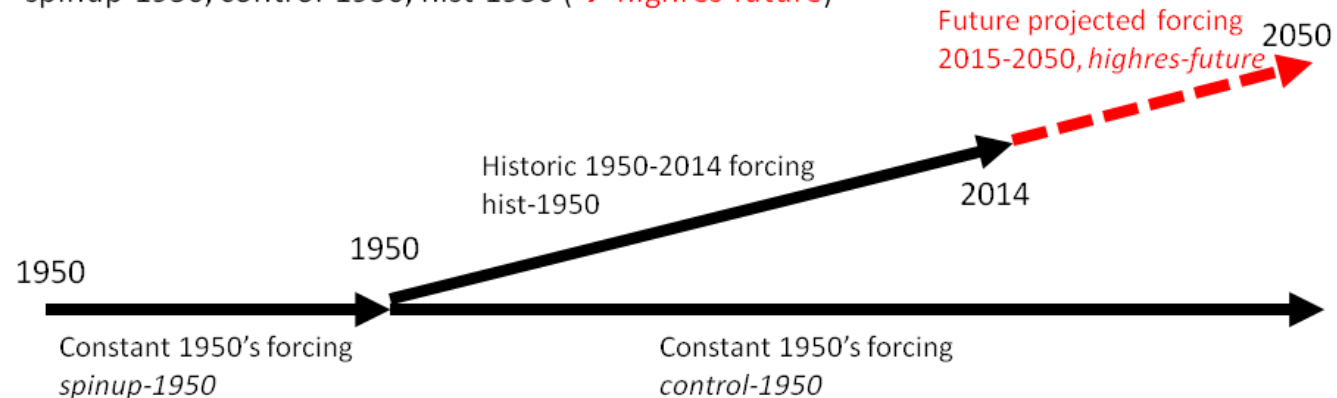


Coupled climate, 1950-2014 (→ 2050)

Forced by constant 1950 and historic forcings (→ projected)

Initial coupled spin-up period ~ 30-50 years from 1950 EN4 ocean climatology

spinup-1950, control-1950, hist-1950 (→ highres-future)



# Models in PRIMavera *running* HighResMIP protocol

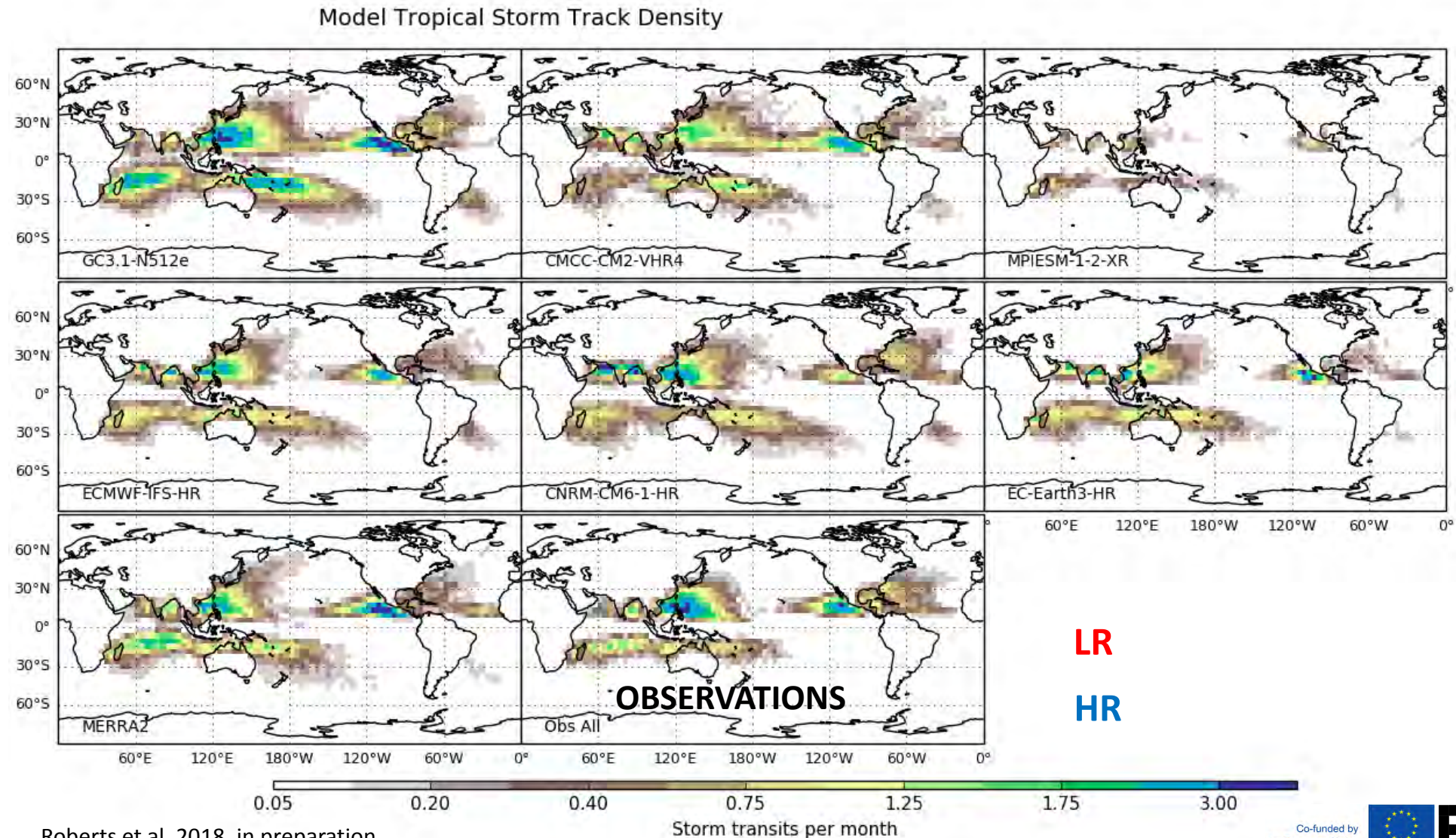
Institution	MOHC, URead, NERC	EC-Earth KNMI,SHMI, BSC, CNR	CERFACS	MPI-M	AWI	CMCC	ECMWF
Model name	HadGEM3 GC3.1	EC-Earth3.3	CNRM-CM6	MPIESM-1-2	AWI-CM 1.0	CMCC-CM2	ECMWF-IFS
Model components	UM NEMO3.6 CICE5.1	IFS cy36r4 NEMO3.6 LIM3	ARPEGE6.3 NEMO3.6 GELATO6.1	ECHAM6.3 MPIOM1.63 MPIOM1.63	ECHAM6.3 FESOM1.4 FESIM1.4	CAM4 NEMO3.6 CICE4.0	IFS cycle43r1 NEMO3.4 LIM2
Atmos dynamical scheme (grid)	Grid point (SISL, lat-long)	Spectral (linear, reduced Gaussian)	Spectral (linear, reduced Gaussian)	Spectral (triangular, Gaussian)	Spectral (triangular, Gaussian)	Grid point (finite volume, lat-long)	Spectral (cubic octohedral, reduced Gaussian)
Atmos grid name	N96 , N216, N512 (L,M,H)	Tl255, Tl511	Tl127, Tl359	T127, T255	T63, T127	1x1, 0.25x0.25	Tco199, Tco399
Atmos mesh spacing 0N	208, 93, 39	78, 39	156, 55	100, 52	200, 100	100, 28	50, 25
Atmos mesh spacing 50N	135, 60, 25	71, 36	142, 50	67, 34	129, 67	64, 18	50, 25
Atmos nominal res (CMIP6)	250, 100, 50	100, 50	250, 50	100, 50	250, 100	100, 25	50, 25
Atmos model levels (top)	85 (85km)	91 (0.01 hPa)	91 (78.4 km)	95 (0.01 hPa)	95 (0.01 hPa)	26 (2 hPa)	91 (0.01 hPa)
Ocean grid name	ORCA	ORCA	ORCA	TP	FESOM (unstructured)	ORCA	ORCA
Ocean res nominal (km)	100, 25, 8 (L,M,H)	100, 25	100, 25	40, 40	50, 25	25, 25	100, 25
Ocean levels	75	75	75	40	47	50	75

6 different atmosphere-only GCMs

7 different coupled GCMs  
(though some common components)

Range of resolutions: from 100km to 20km  
... and further to sub-10km

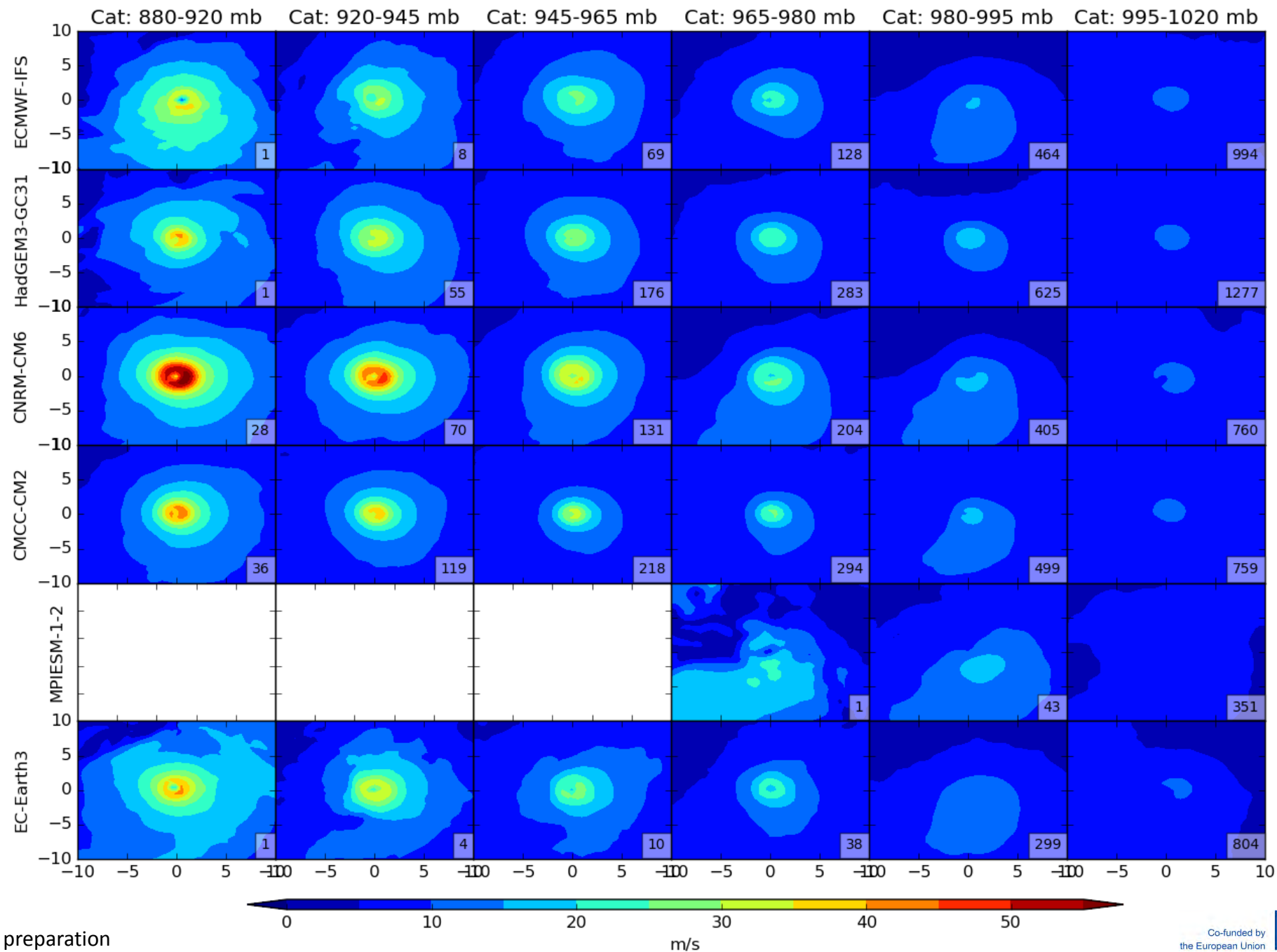
# Tropical Cyclone track density: 65 year climatologies (storm transits per month per 4 degree unit area)



High resolution

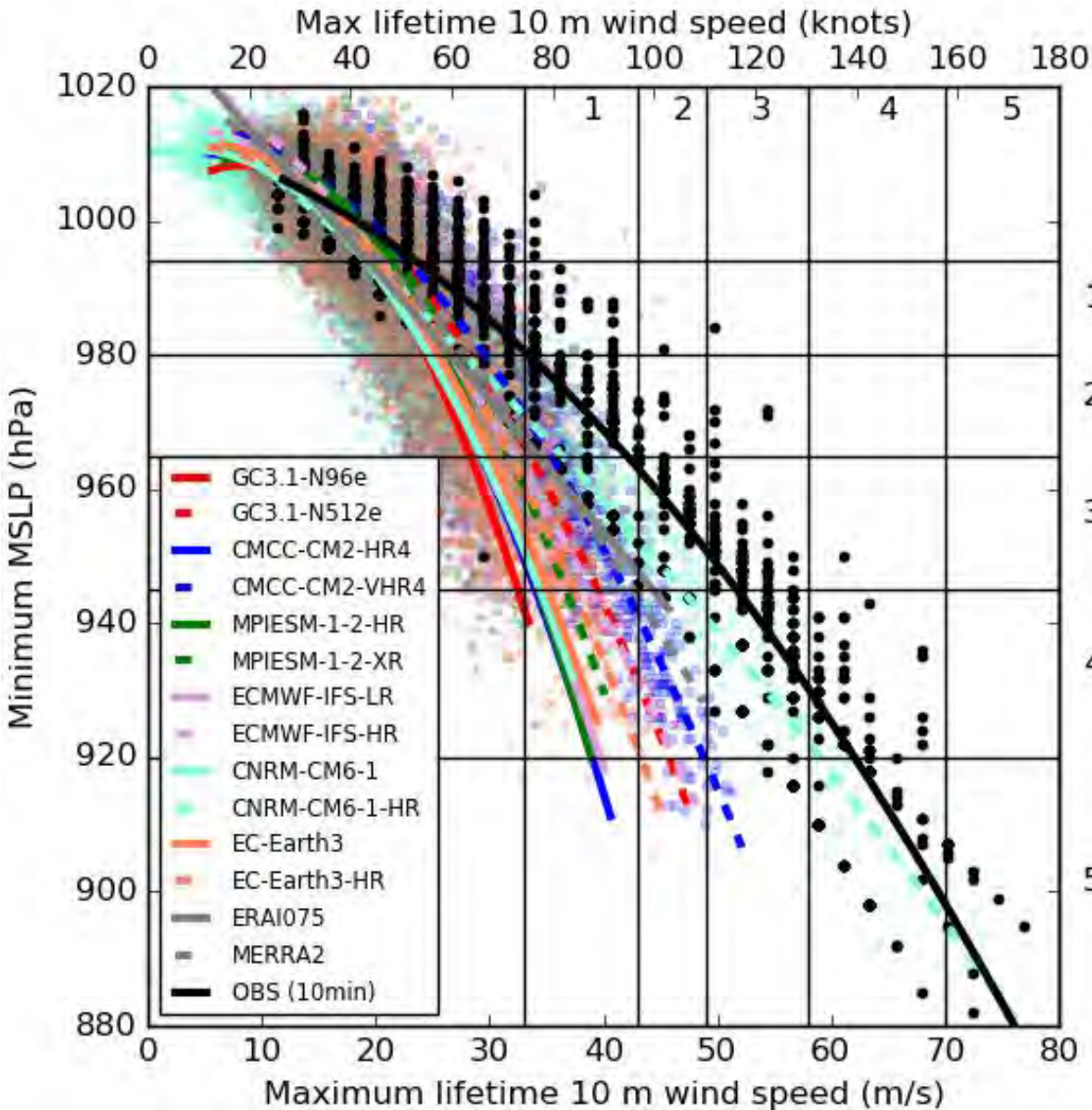
Composite HR storms for: near\_surface\_windspeed

Low resolution



# TC intensity using MSLP-10m wind

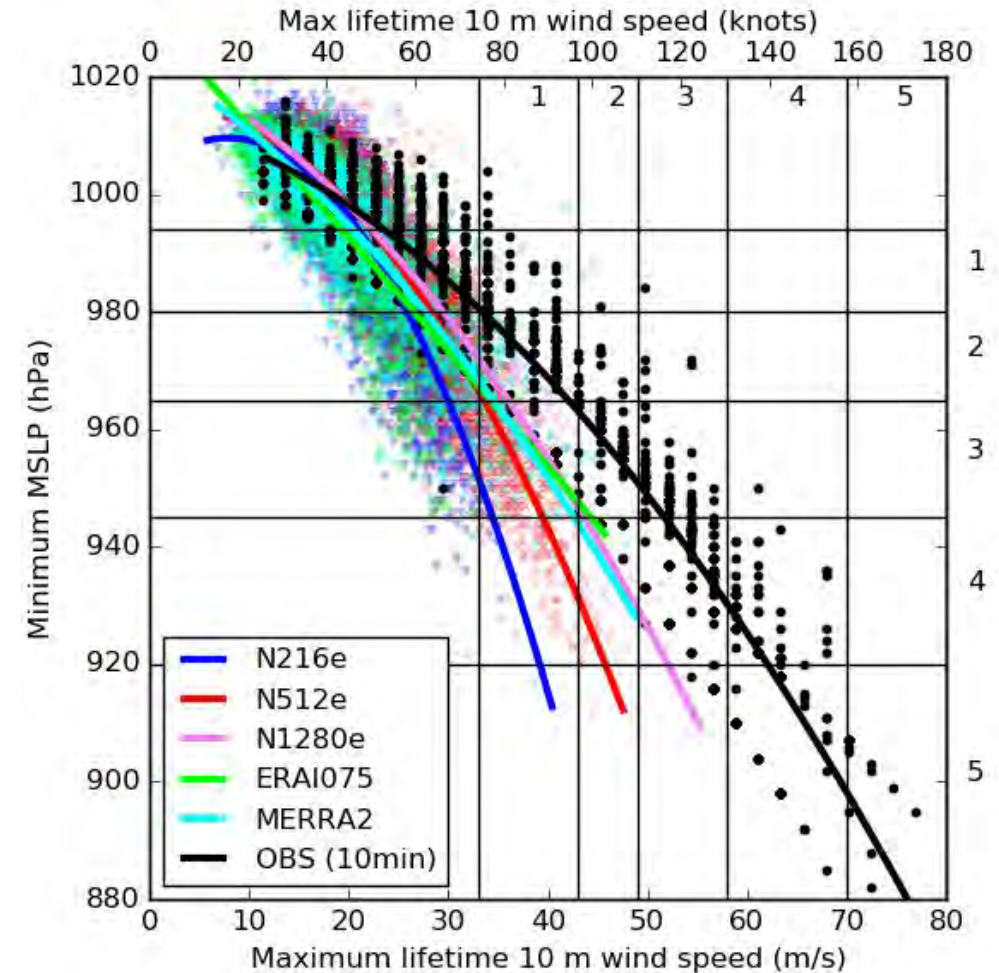
(instantaneous 6 hourly, not max/min over 6 hours)



Continuous  
lines are  
coarser  
GCMs

Dashed  
lines are  
higher  
resolution  
GCMs

HadGEM3, from 100 to 10km resolution



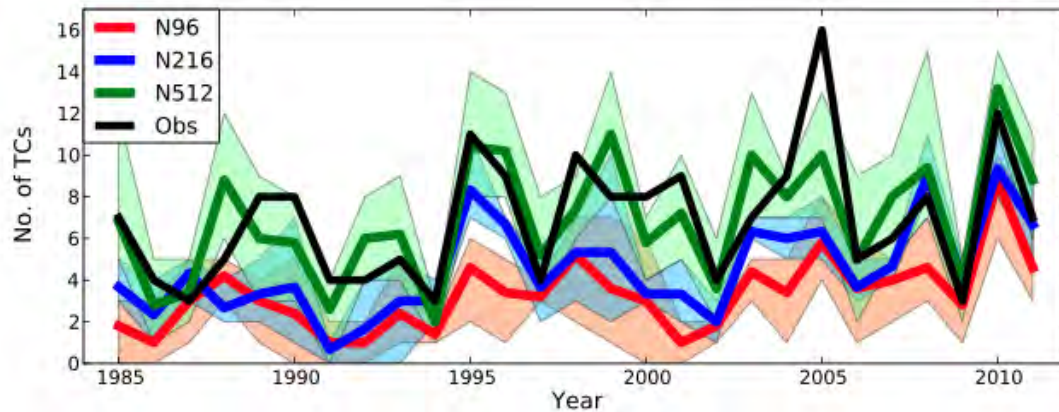
# Interannual TC frequency correlation with observations (all/hurr) - 1 member

One of the most important results in the CLIVAR HWG experiment was this: **skill at representing interannual variability improves with model resolution.**

→ Key to seasonal prediction of hurricanes (and typhoons)

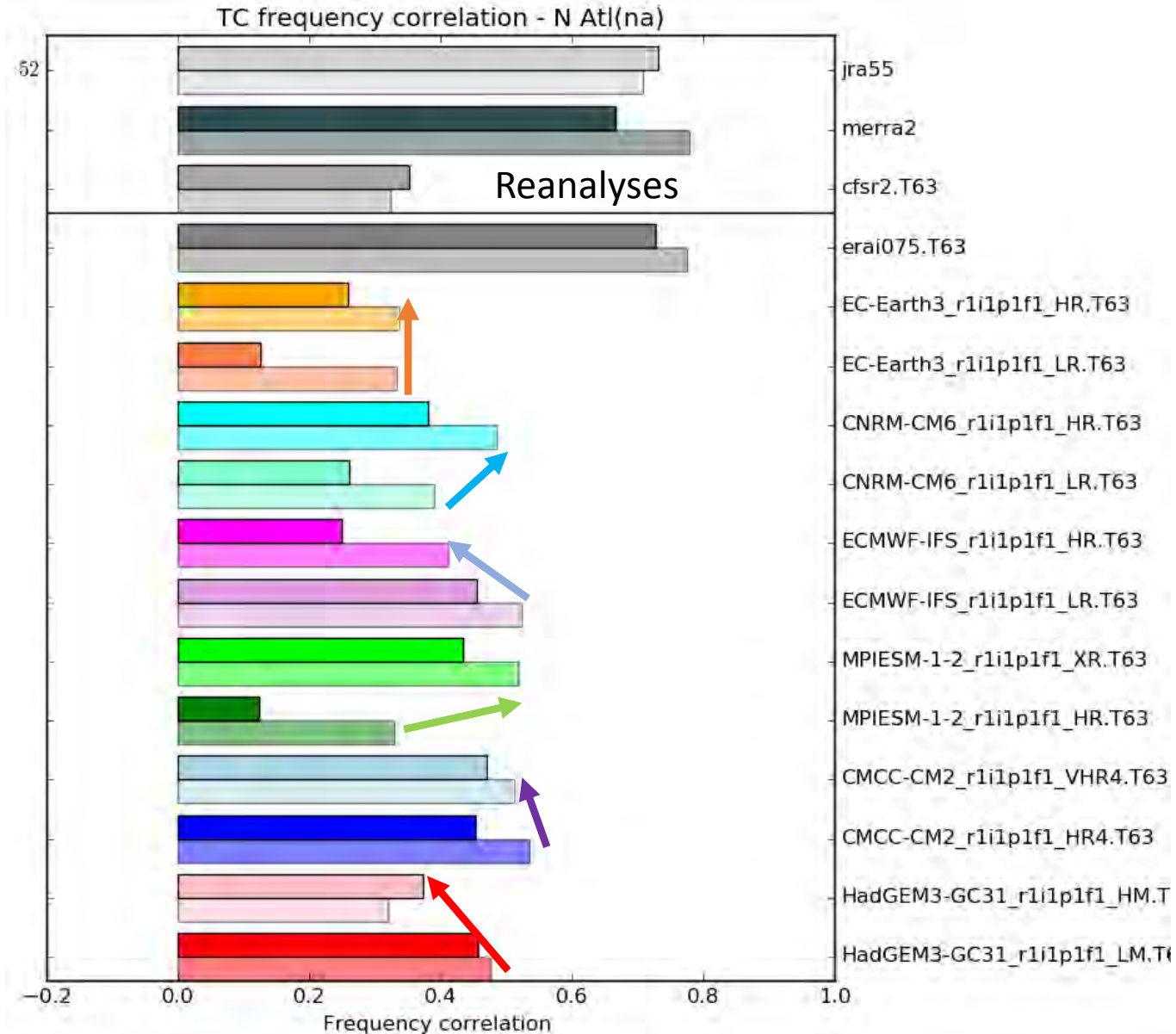
In 2015, as part of our work in the *US CLIVAR Hurricane Working Group* using our **2012 PRACE-UPSCALE** data:

TC frequency, track density and interannual



Roberts et al. 2015. Journal of Climate

Previously also shown in Zhao et al. (2010) and Strachan et al. (2011)



Roberts et al. 2018, in preparation

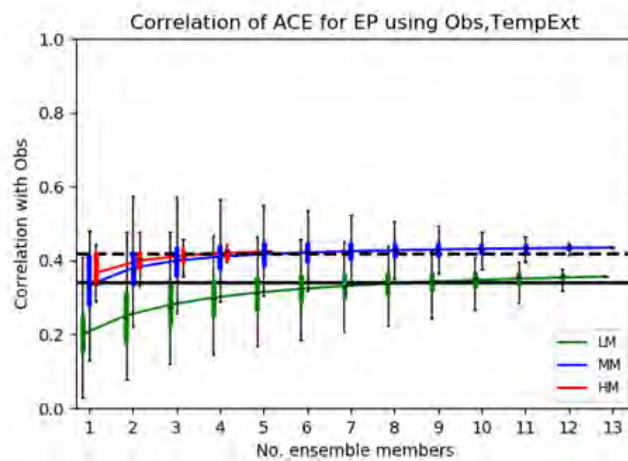
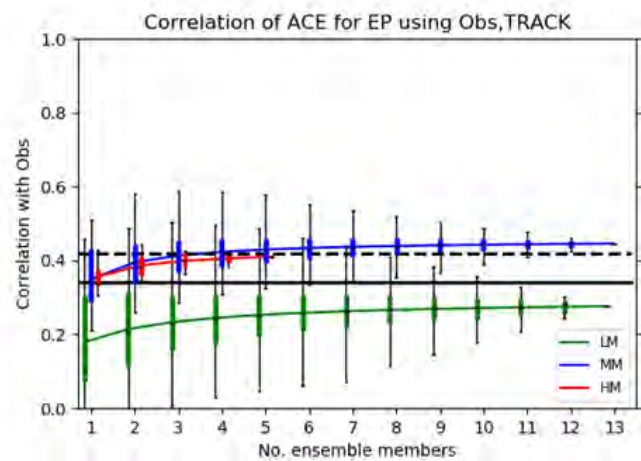
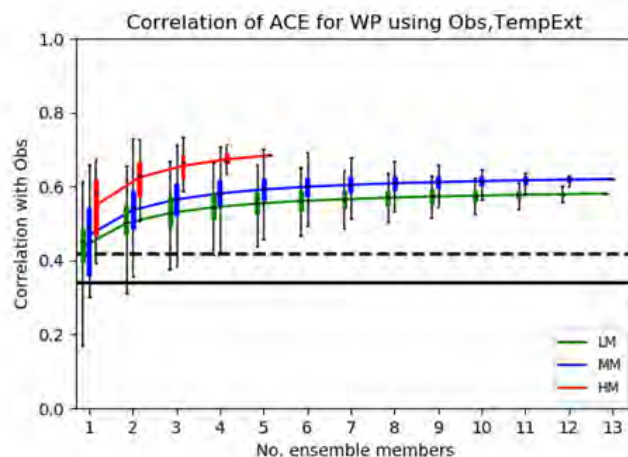
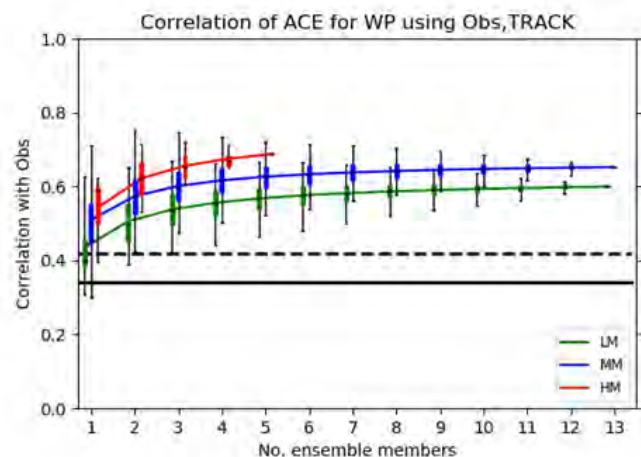
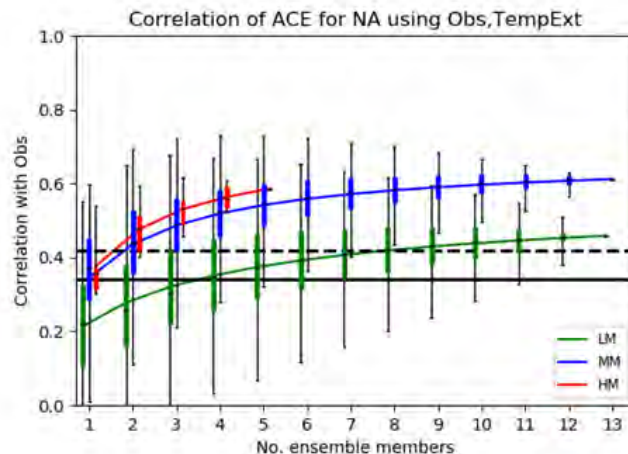
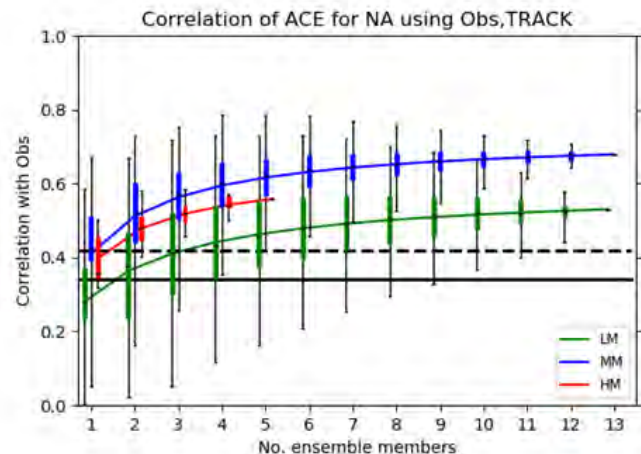
Is using single ensemble members per GCM enough to robustly represent interannual variability?

Multiple GCM resolutions of ensembles, 2 tracking algorithms

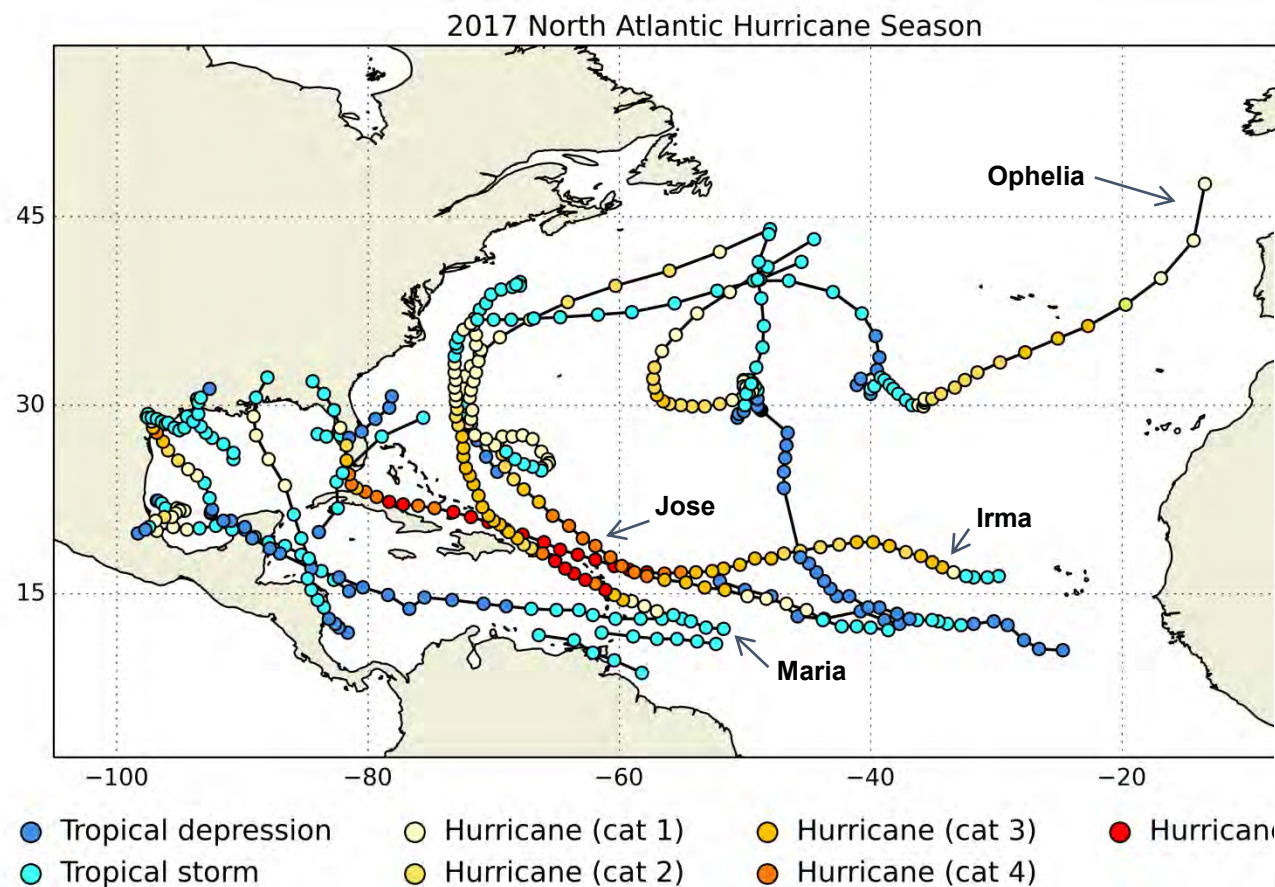
At least **6 ensemble members needed** in the North Atlantic to understand skill in simulating interannual variability

**3-4 ensemble members seem sufficient** in the West Pacific.

We do have a heterogeneous ensemble in PRIMAVERA, but also small ensembles of each GCM. → need to revisit IV



# Motivation II: a changing risk from TCs

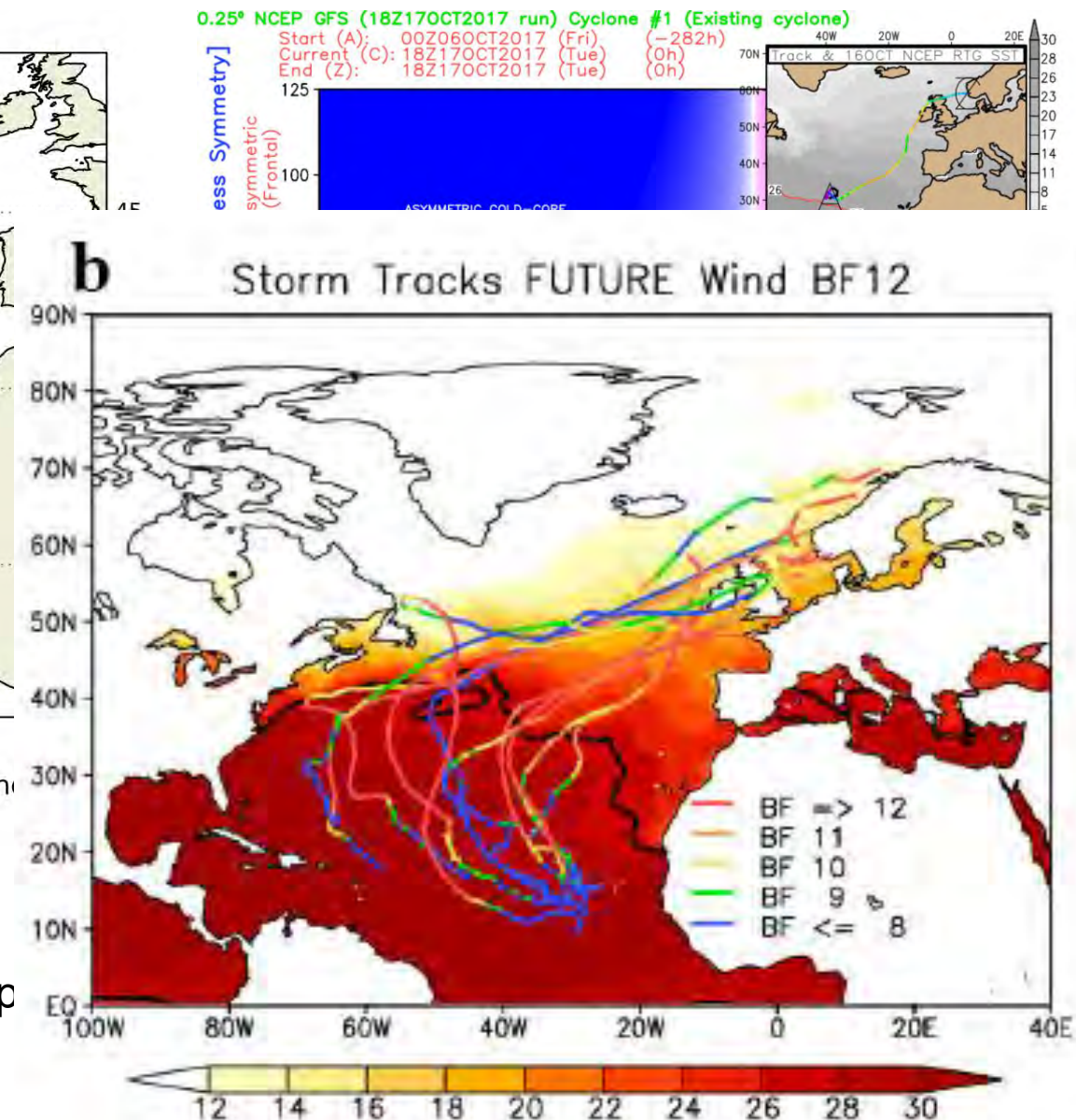


**Total storms: 17**

- 17 tropical storms (39+ mph)
- 10 hurricanes (74+ mph)
- 6 major hurricanes (111+ mph)

Accumulated Cyclone Energy (ACE) index = 226

Thanks to Jo Camp



# Track density

# Genesis density

ERA-Interim

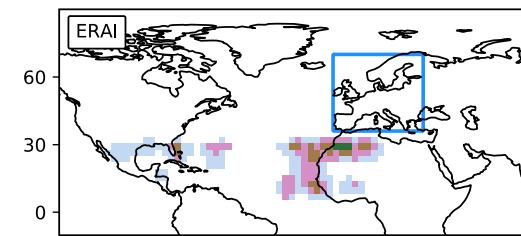
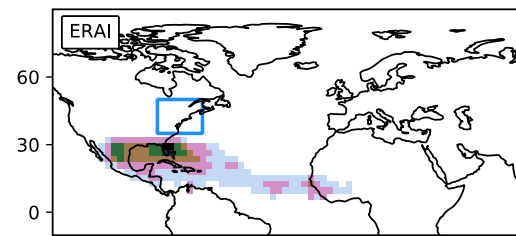
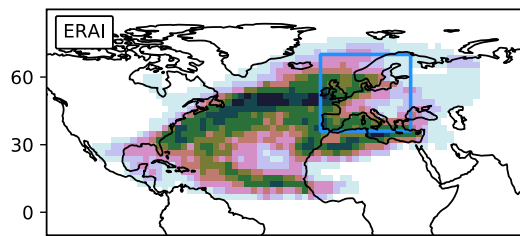
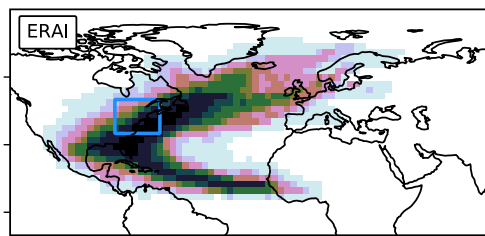
northeastern US

Europe

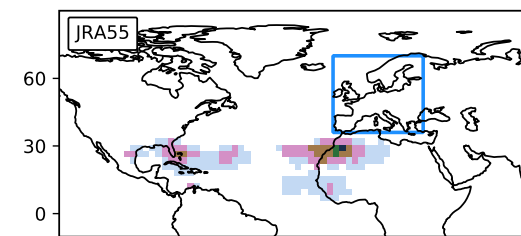
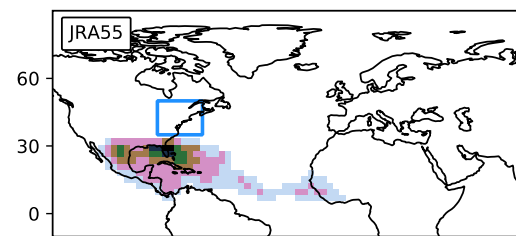
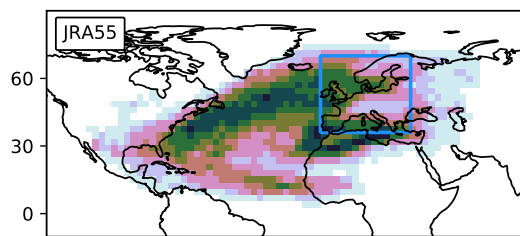
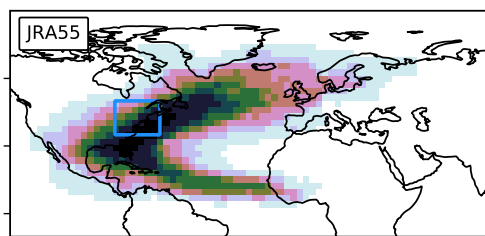
northeastern US

Europe

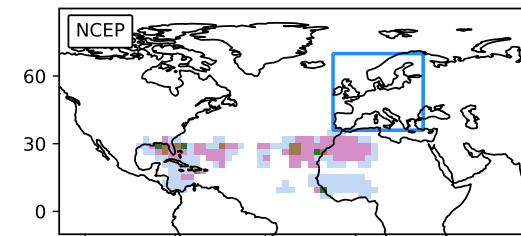
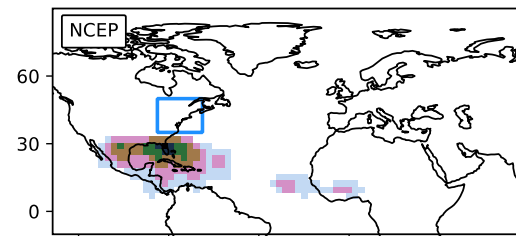
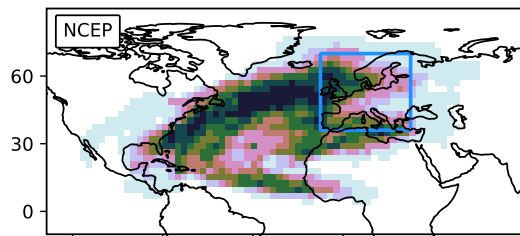
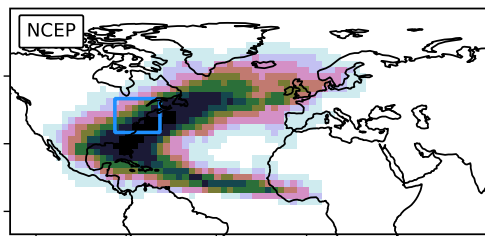
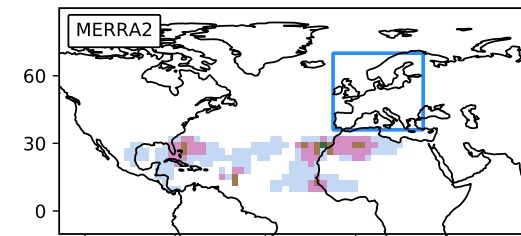
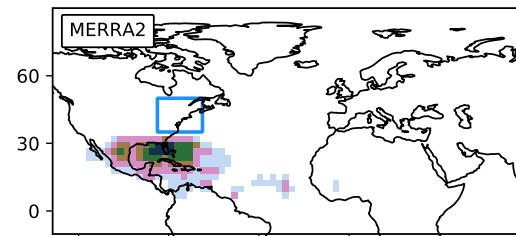
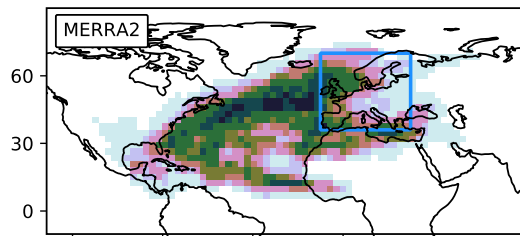
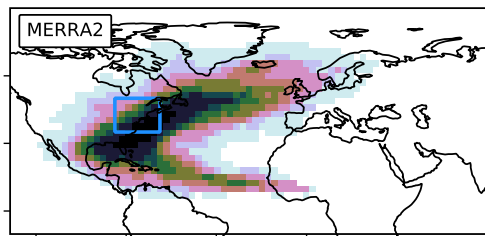
JRA-55



MERRA2



NCEP-CFSR

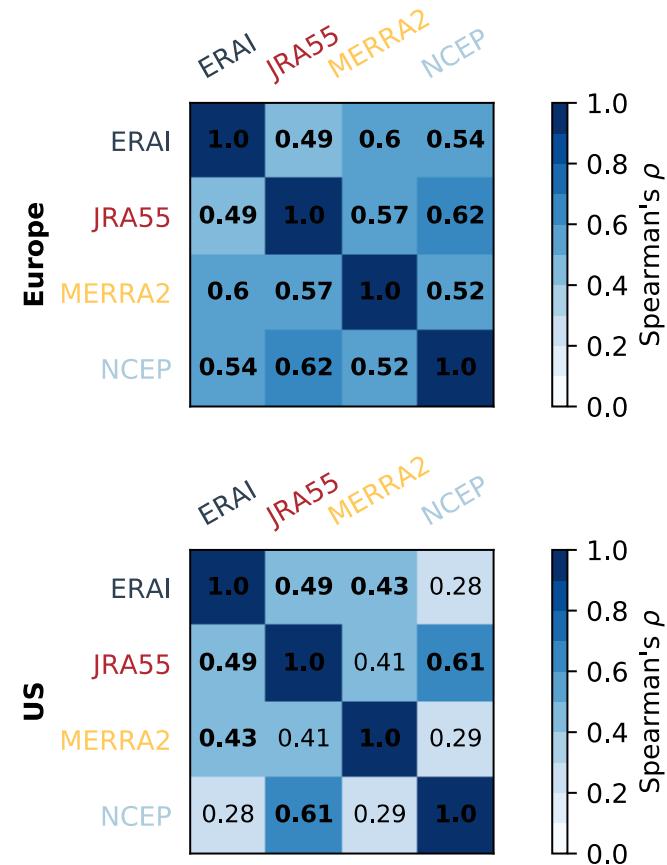
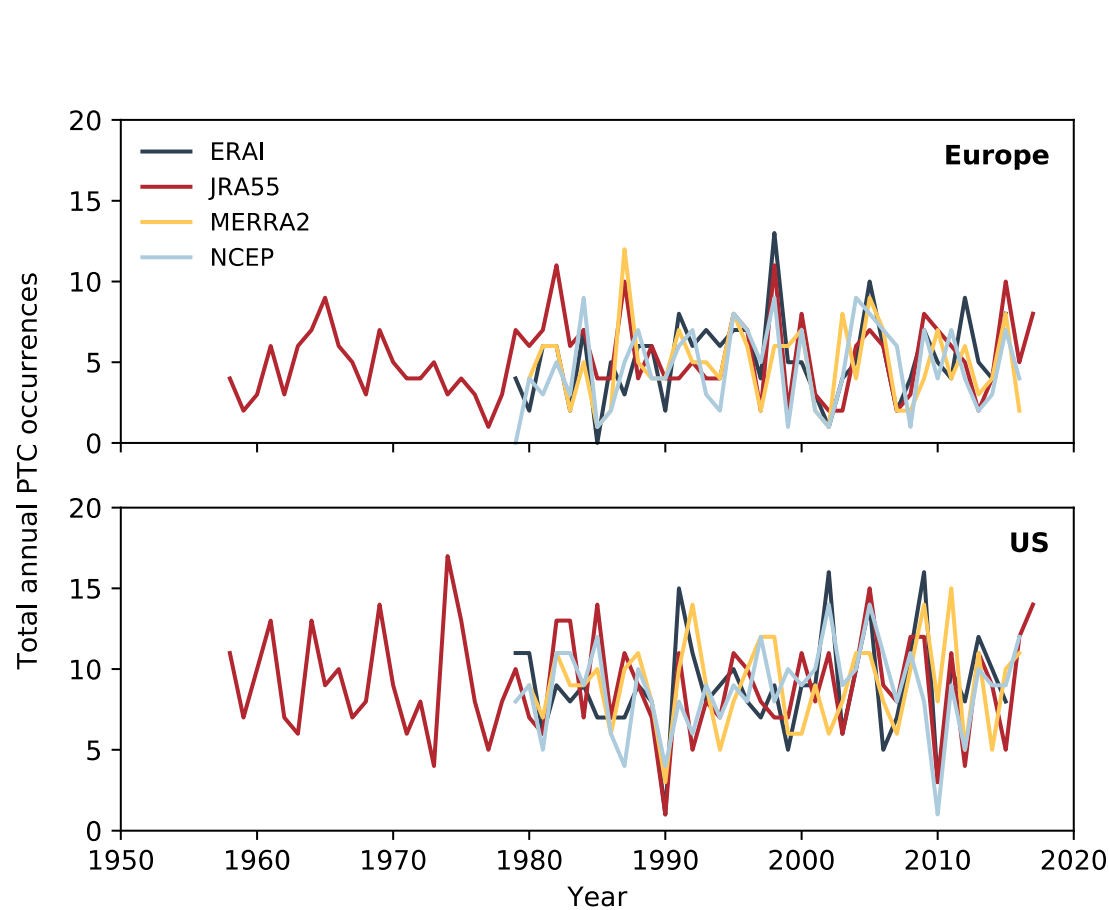


0.0 0.1 0.3 0.5 1.0 1.5 2.0 3.0 4.0 8.0 16.0  
cyclones per month per unit area

ANNUAL MEAN

0.0 0.1 0.2 0.5 1.0 2.0 3.0 4.0  
cyclones per month per unit area

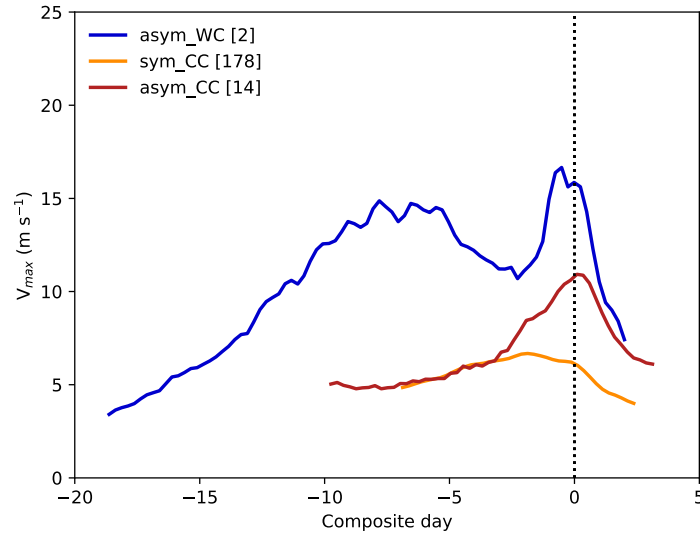
# Historical variability of North Atlantic post-tropical cyclones



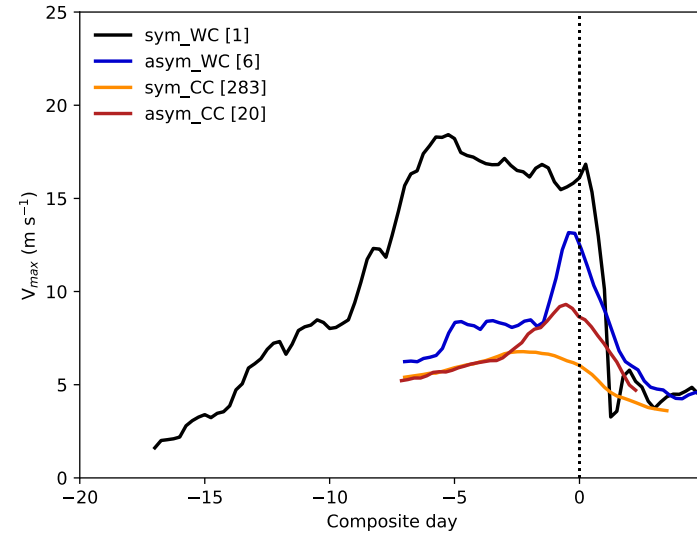
**\*bold = significant at 99% level**

# Composite post-tropical cyclone lifecycles

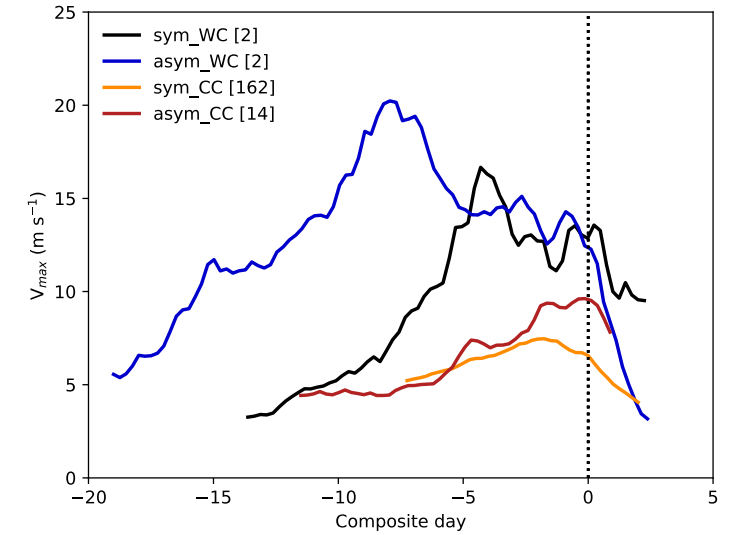
ERA-Interim, EUROPE, 1979-2015



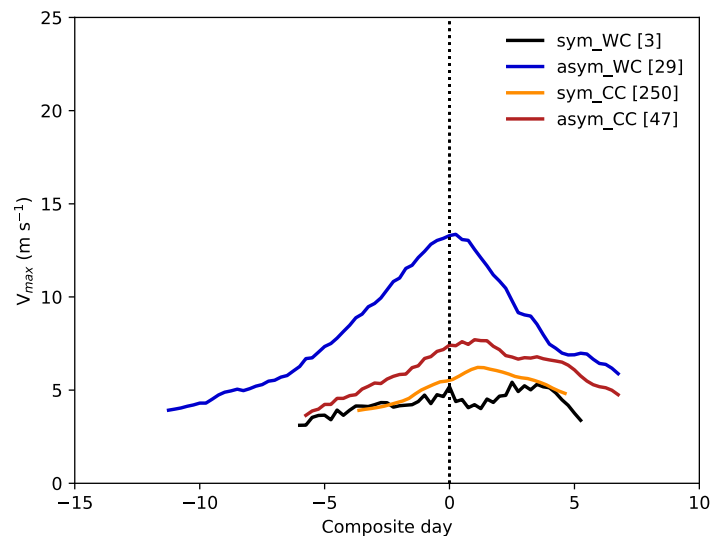
JRA55, EUROPE, 1958-2017



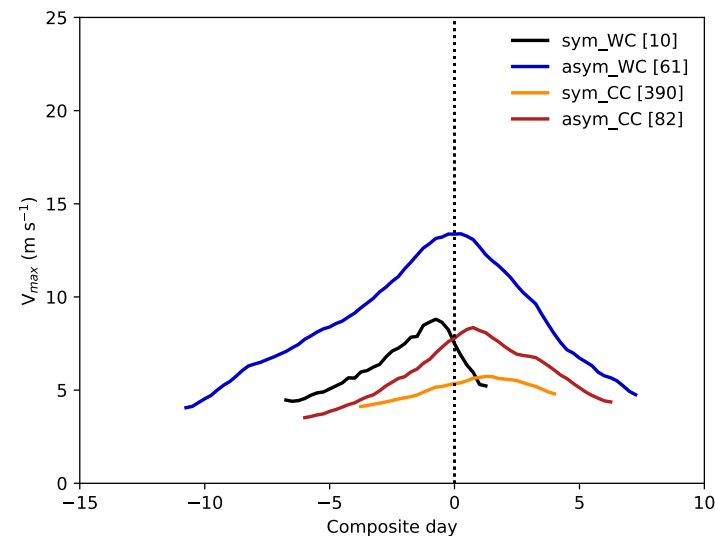
MERRA2, EUROPE, 1980-2016



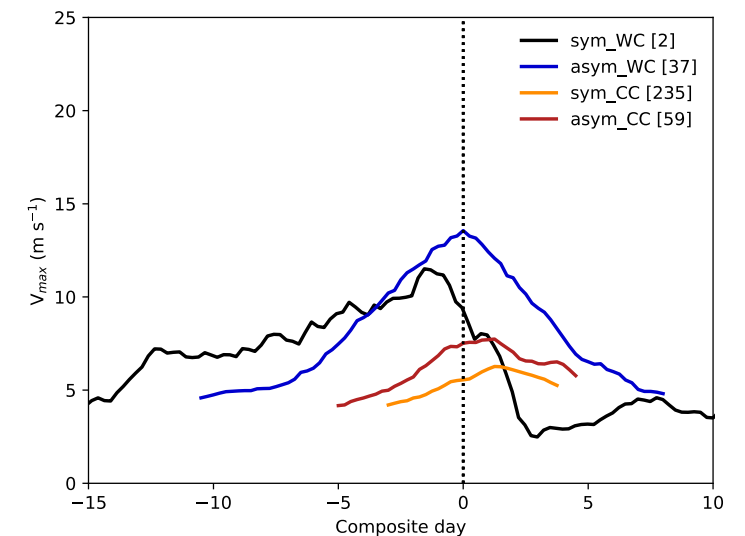
ERA-Interim, US, 1979-2015



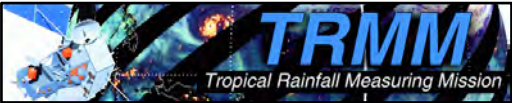
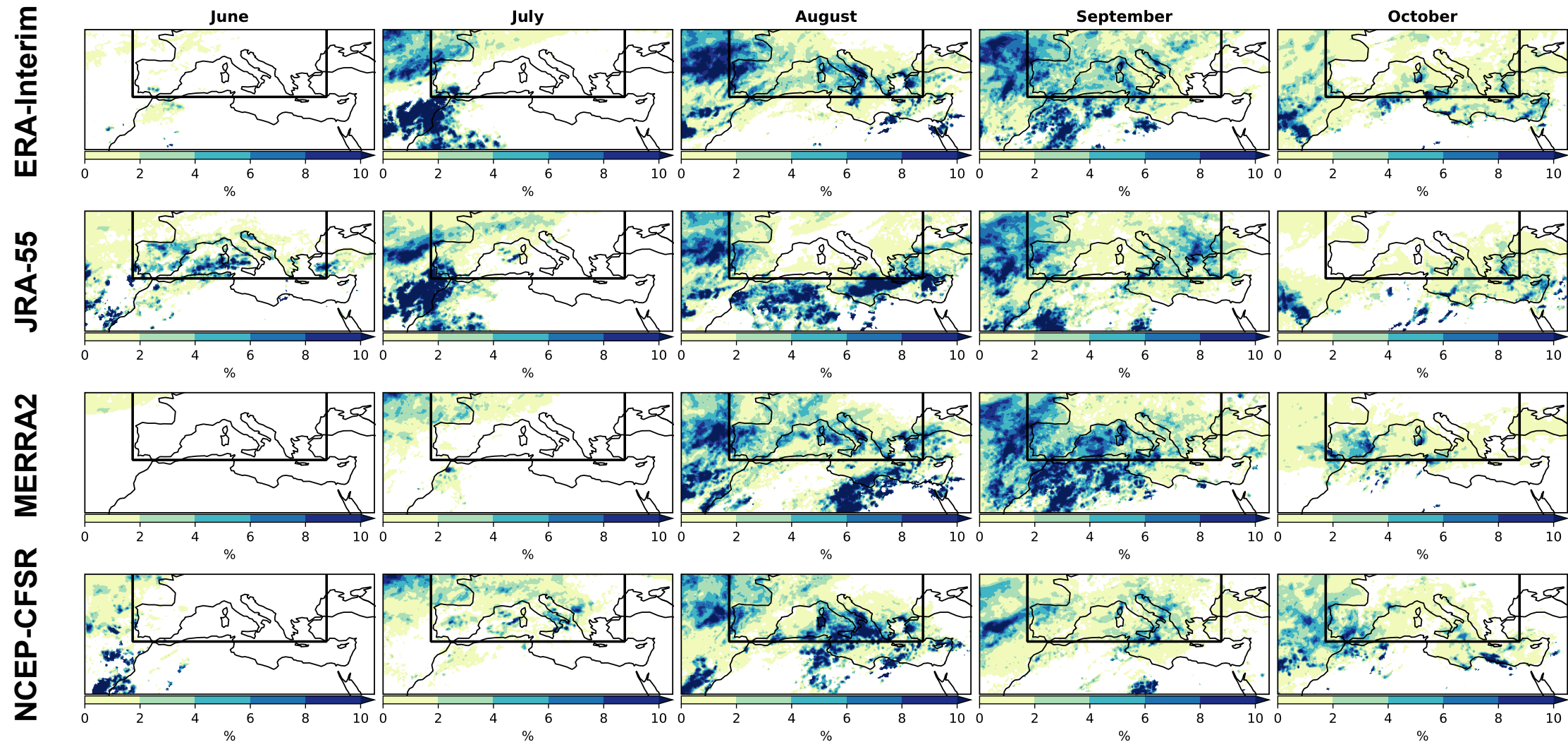
JRA55, US, 1958-2017



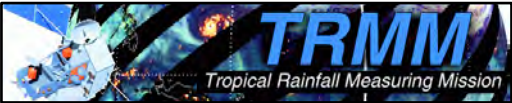
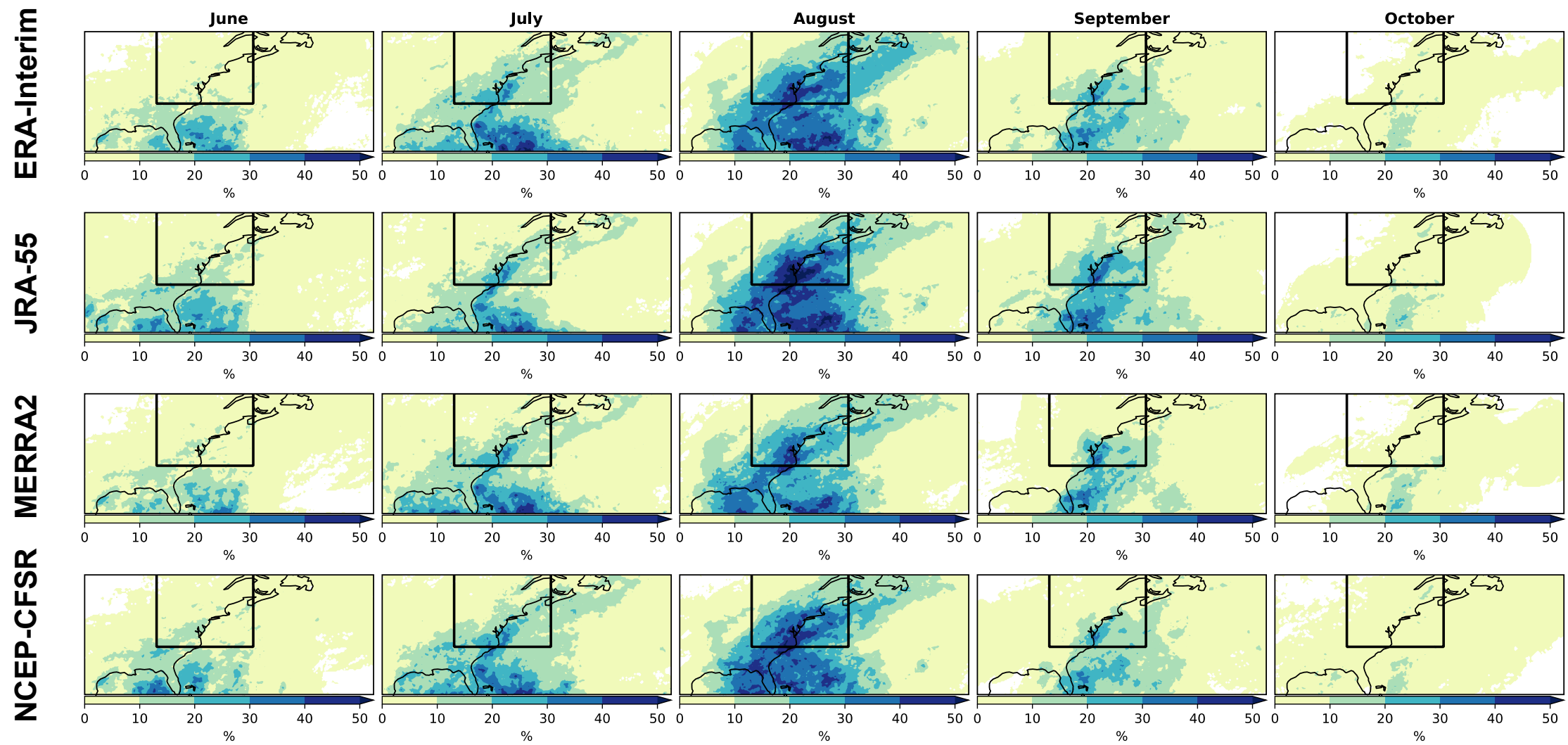
MERRA2, US, 1980-2016



# Post-tropical cyclone-associated precipitation – Europe



# Post-tropical cyclone-associated precipitation – northeast US



# Summary

- Tropical Cyclones emerge in high-resolution GCMs:
  - The good and the bad:
    - 50-20km resolution: credible representations of track density and interannual variability
    - Structures are credible at ~20km, but TCs still too large in most GCMs.
    - Intensity at ~20km still not sufficient to capture all CAT4,5 in most GCMs (but there are notable exceptions with full spectrum represented)
  - In this context it is extremely important that we are working as a community, under **an agreed protocol, HighResMIP**. Progress in understanding from that decision. However:
    - We need ensemble size of at least 5 to robustly represent interannual variability
  - TC-ENSO relationship credibly represented in historical simulations
  - Poleward shift of TCs seen in climate change projections by GCMs capable of resolving TCs
- We are working towards sub-10km GCMs: expect better skill in terms of intensity.
- Much work left to do on post-tropical cyclones, extra-tropical transition, structures, etc.

# HRCM and TC research: who is doing what?

Topic	Sub-topic	People
Storm tracks, variability		M Roberts, K Hodges
TCs simulation	Role of resolution Role of Stochastic Physics	M Roberts, PL Vidale PL Vidale, K Hodges, M Robersts, ECMWF group
Post-Tropical Cyclones		A Baker
TC-centred hydrological cycle		B Vanniere
TC energetics		B Harris
Decadal Variability		P Loizou
Impacts	Climatologies of precipitation(and mositure transports) caused by TCs	L Guo, A Franco Diaz

Co-funded by  
the European Union



**PRIMAVERA**