Airmass analysis of the processes driving the progression of the 2016 Indian summer monsoon

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- Monsoon flow is from SW **<u>but</u>** ISM propagation is SE-NW:
- erosion of wedge-shaped dry air layer by moist low-level SW-ly flow underneath
- Link between heavy monsoon rainfall and the passage of mid-latitude troughs, see Martius et al. (2013) and Vellore et al.(2016)
- Link between dry advection over India and:
- blocking over W Asia (Krishnamurty et al., 2010)
- anomalies in the subtropical jet and advected PV (Fletcher et al., 2018)



- Is ISM progression, along with the erosion of the dry layer, a steady process or does it happen in "bursts", like for the Australian monsoon? (Berry and Reeder, 2016)
- How is it linked with the extratropical circulation?
- > What is the importance of diabatic processes in the removal/moistening from below of the dry layer?

This study (Volonté et al., submitted):

- Regional analysis: 17km UKMO global operational forecast
- Mesoscale analysis: 4.4km convection-permitting UKMO LAM (INCOMPASS seasonal run)







RH at 760 hPa:

- Net reduction in dry-air area between 10 and 25 June: expansion and deepening of moist monsoon flow, particularly over NW India
- Maximum southward expansion of dry air on 15 June
- Cyclonic circulation in the Arabian sea separating dry and moist air on 20 and 25 June













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National Centre for Atmospheric Science What causes the southward advection of dry air over the Arabian sea around 15 June?

T & gph at 625 hPa + upper-level jet:

- Passage of mid-latitude trough on 13 June
- Formation of cyclonic circulation over the Arabian Sea on 16 June
- Are these two features linked?







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PV at 700 hPa:

- High-PV associated with the trough is advected towards SE, after its passage on 13 June
- The cross section shows the presence of a descending PV streamer
- This streamer helps the formation of a cyclonic system when on top of lowlevel high-PV air ("PV tower", see Čampa and Wernli, 2012)

(a) 13 June - 06UTC (b) 14 June - 18UTC 114 24 2 0.5 (c) cross section AB on 14 June - 18UTC 1.7 <u>د 600</u> 700 800 900 A 20 m/s — - 600 2400 1200 1500 Distance (km)

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0.<u>9</u>

0.5



- Lagrangian trajectories help us understanding origin and path of different air masses
- Core of moist air:
 "classic" Somali Jet path towards India
- Core of dry air:
 partly descending and advected from WNW
- 10-15 June: more interaction between dry and moist flows as the dry incursion flows over the Arabian Sea
- 20-25 June:

Bigger separation between the two air masses, no penetration of the dry air beyond Hindukush range



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hi-res LAM simulation: a closer look at the interaction between the two different air masses over the northern stretch of the Arabian Sea (e.g. on 15 June)

- 2-layer structure confirmed with dry NW-ly advection flowing above moist SW-ly monsoon flow



4.4km convection-permitting UKMO LAM (INCOMPASS seasonal run)







- Cross sections show:
- 2 layers: NW-ly dry above, SW-ly moist below
- Sharp boundary between the 2 air masses at 850 hPa
- Occurrence of deep convection prevented by the dry layer

(c) AB cross section



(d) CD cross section



4.4km convection-permitting UKMO LAM (INCOMPASS seasonal run)







- Thermodynamic profiles show : \geq
- sharp transition between the two air masses and the role of dry air in suppressing convection (Profile 1)
- lower troposphere still very dry over the Indus valley ٠ (Profile 2)



(a) 700 PPa







(b) 925 hPa

56 Î

LAM trajectories (e.g. on 10 June): the cores of the two air masses travel along similar paths over the Arabian Sea and then converge while moving inland



4.4km convection-permitting UKMO LAM (INCOMPASS seasonal run)









Trajectory profiles reveal that:

- The moist airstream receives a substantial part of its moisture from the Arabian Sea
- Strong mixing occurs over land when the two airstreams interact
- Moist processes, along with mixing, considerably change the thermodynamic properties of these airstreams



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Summary:

- Progression of the ISM toward NW India is a non-steady process (2-layer model: "tug-of-war" between SW-ly low-level monsoon flow and NW-ly descending dry-air incursion)
- Balance between these airstreams can be influenced by synoptic-scale dynamics at higher latitudes
- Cores of dry and moist airstreams can interact very closely during the onset and progression of the monsoon,
- Importance of diabatic processes, particularly over the Arabian Sea

Next steps:

- Quantitative diagnostics evaluating equatorial and extratropical influences on dry-air incursion strength and associated monsoon "bursts" (e.g. using reanalysis to compute vorticity fluxes towards the Indian subcontinent, see Narsey et al. (2017) for Australian monsoon)
- Analysis of moisture and heat surface fluxes associated with the evolution of the airstreams to understand which diabatic processes are at play
- Lagrangian analysis of moisture sources to assess the role local and remote evaporation/precipitation in the transport of moisture along the airstreams towards the monsoon core region.





